



UNIVERSITY OF GOTHENBURG

The Arctic Ocean Carbon Cycle in a Changing Environment

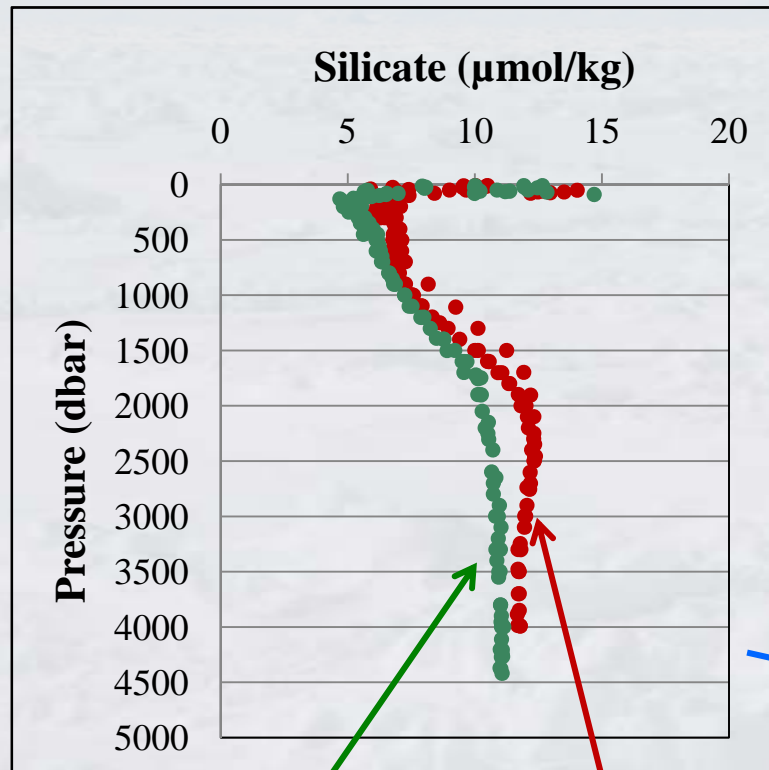
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Faculty of Science

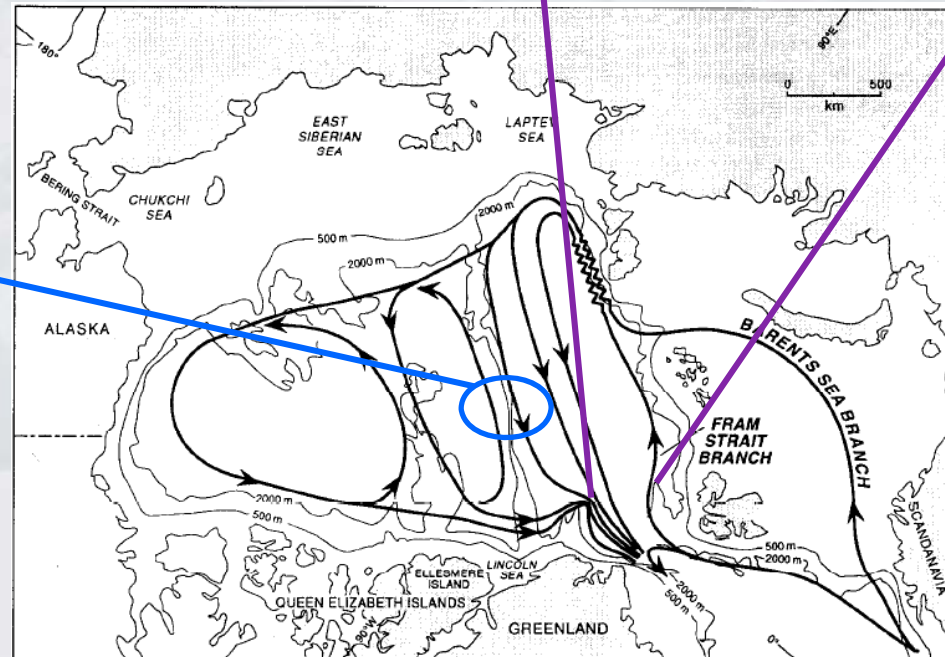
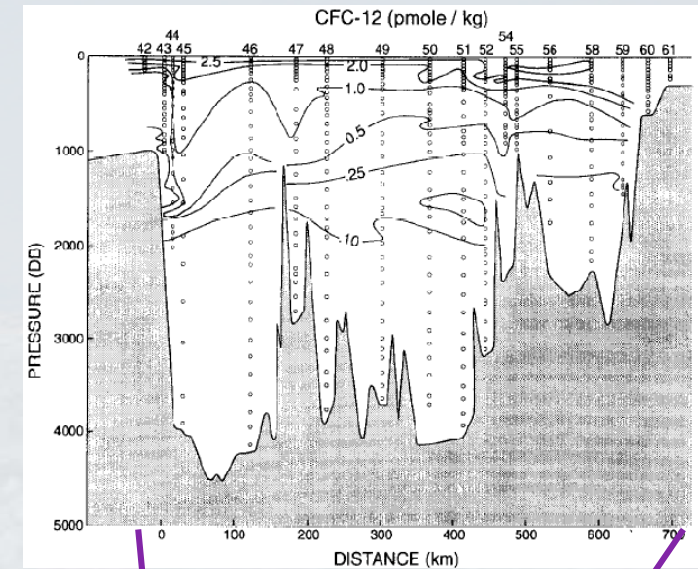
Arctic Ocean circulation

Intermediate & deep waters follows the topography counter-clockwise.



Amundsen
Basin

Makarov Basin

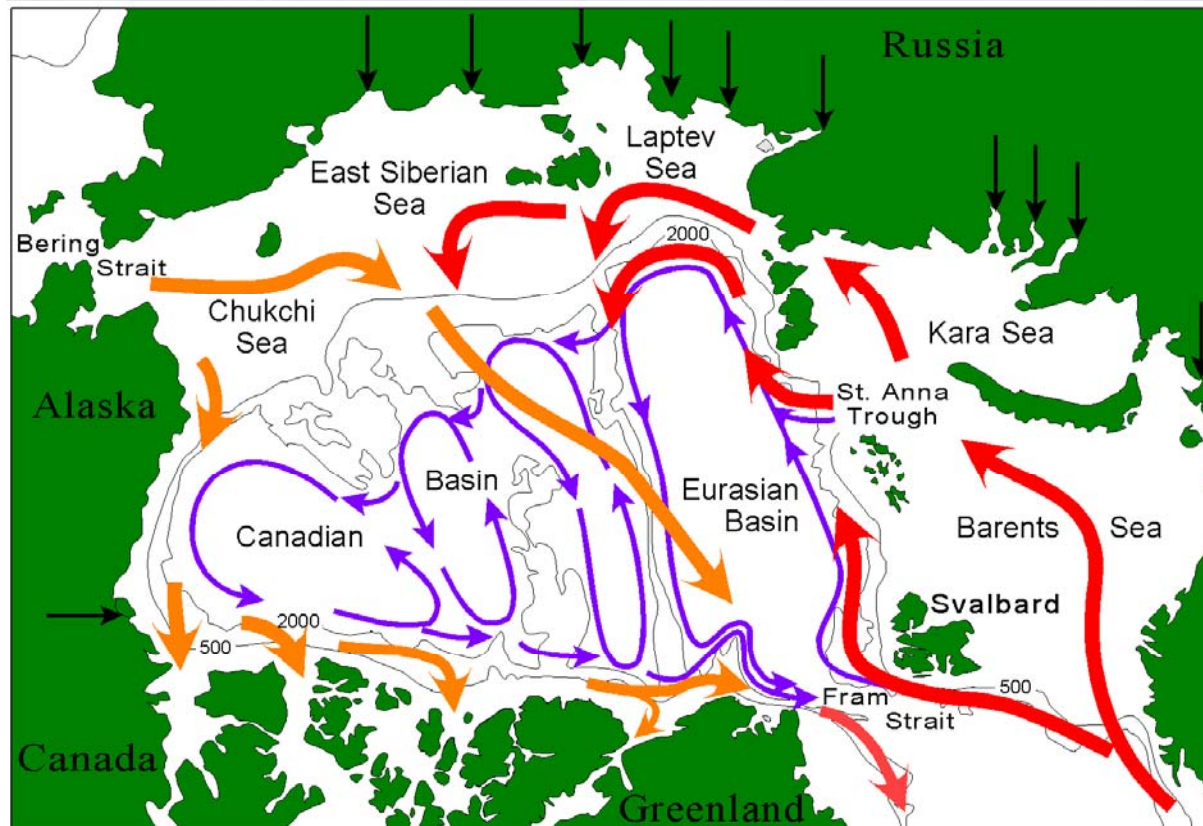


Rudels et al.

Arctic Ocean circulation

Surface water circulation

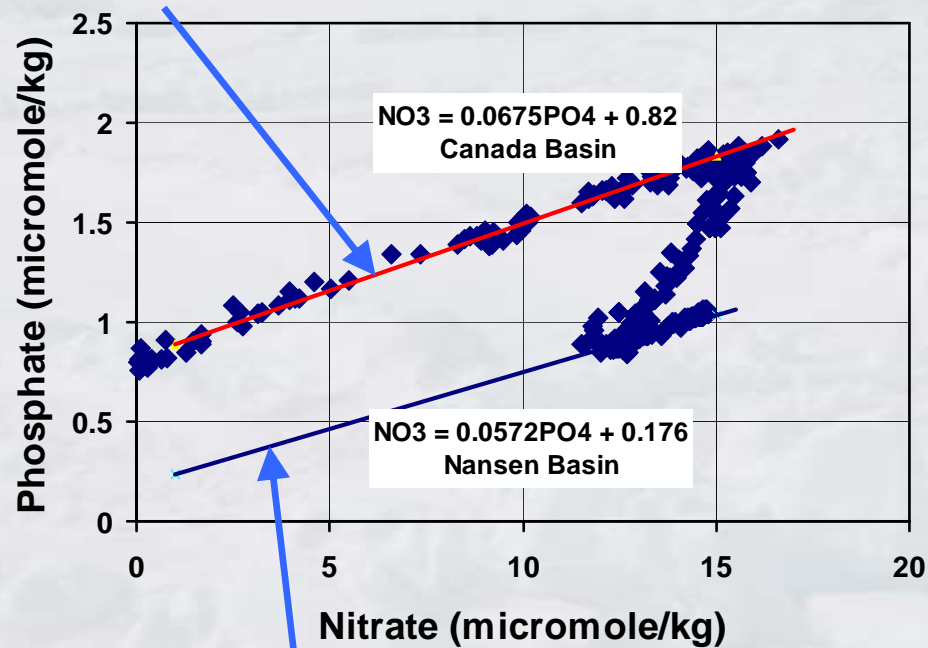
Much of the Pacific and Atlantic waters enter over the shelves and adds to the surface waters of the central basins, but what are their flow paths?



Chemical signatures solve this question

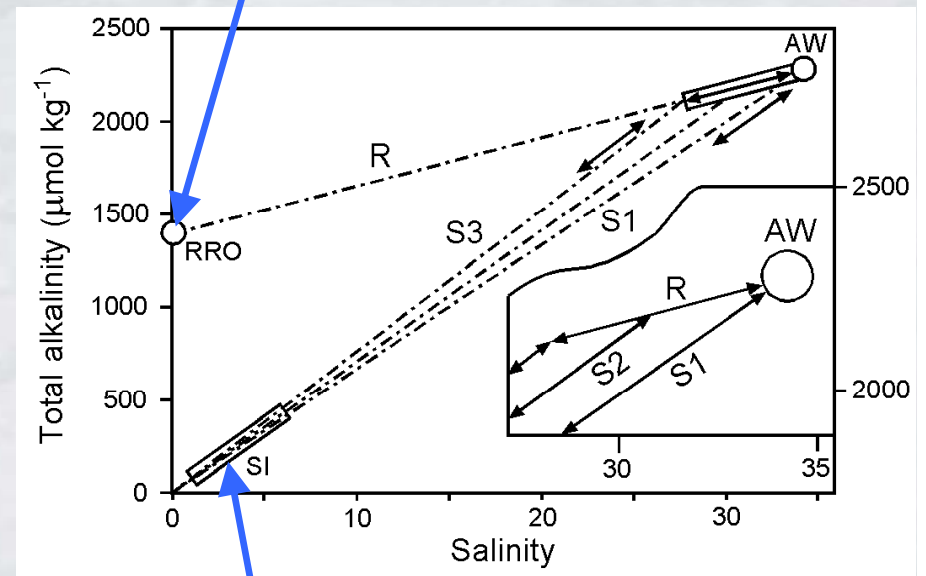
Tracing surface water sources by applying relationships of phosphate - nitrate, and total alkalinity - salinity.
($\delta^{18}\text{O} - S$)

Arctic Ocean Phosphate-Nitrate Relationship
Pacific Water, $S \sim 32$



Atlantic Water

River Runoff



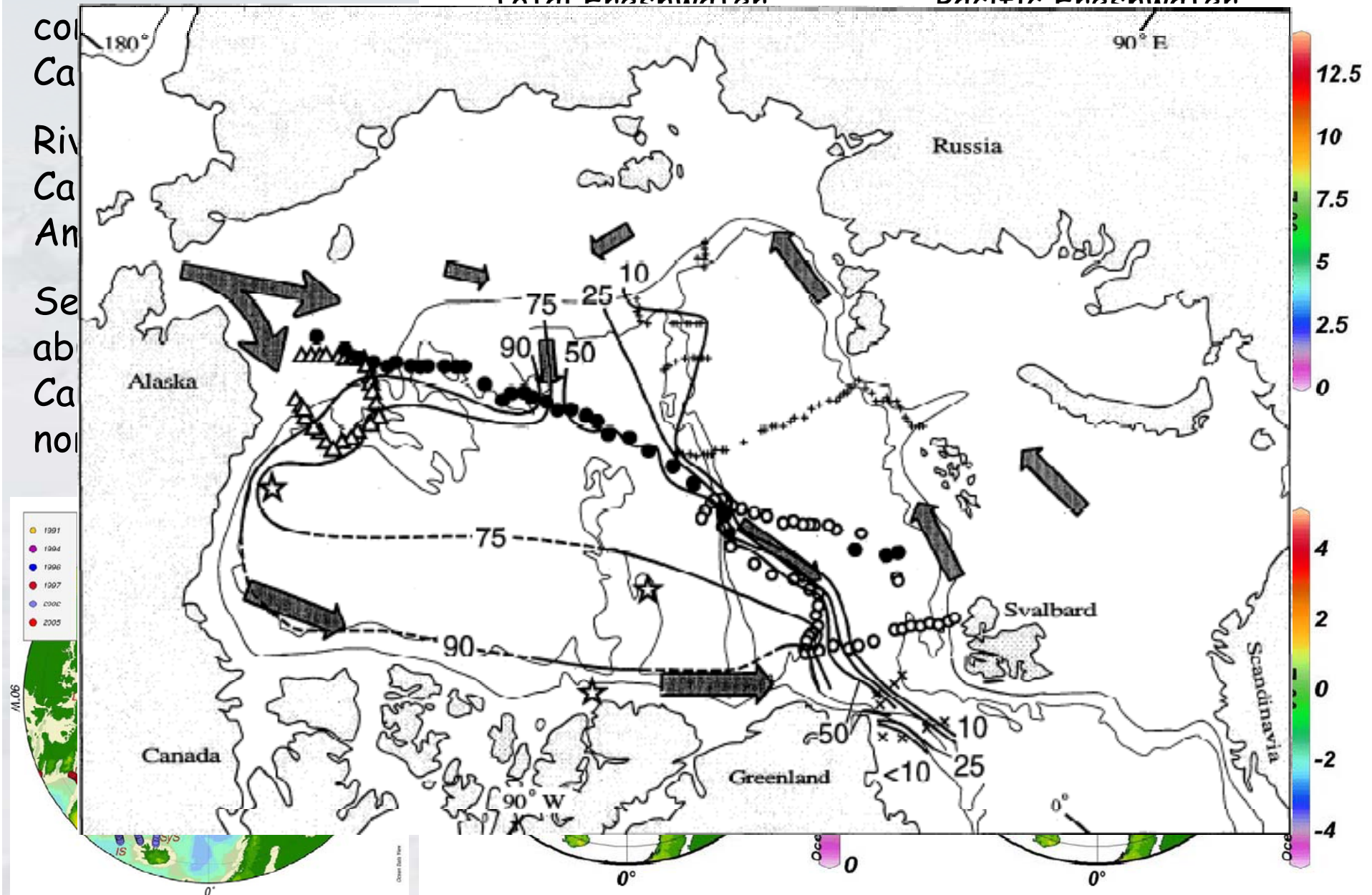
Sea Ice Melt

Jones et al.

Pacific freshwater is

Total Freshwater

Pacific Freshwater

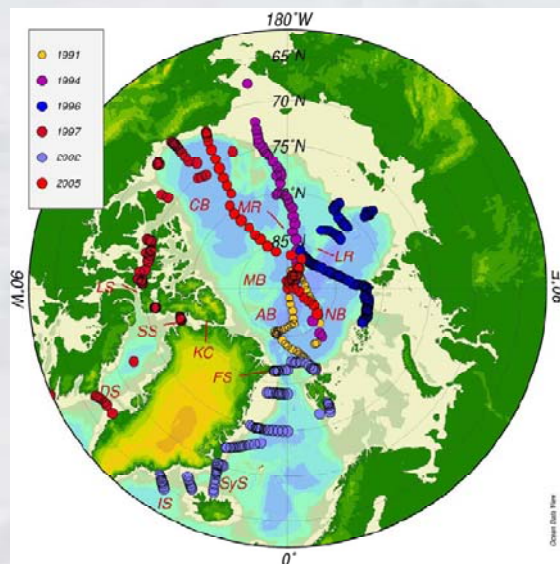


Except for circulation this approach also gives the source waters

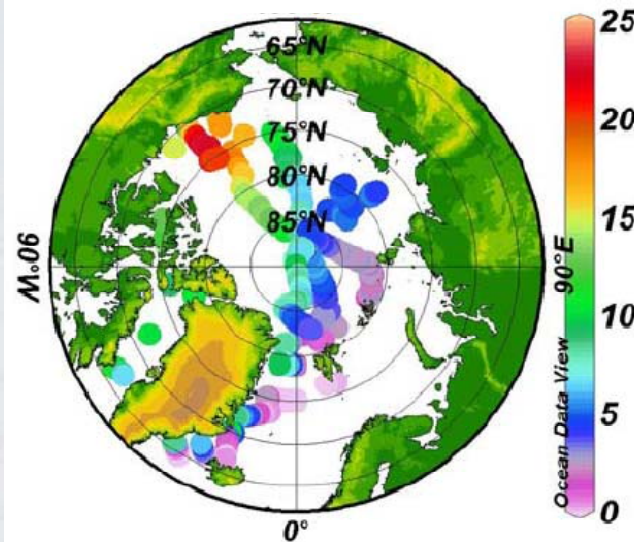
Pacific freshwater is confined to the Canada Basin.

River runoff to the Canadian and Amundsen Basins.

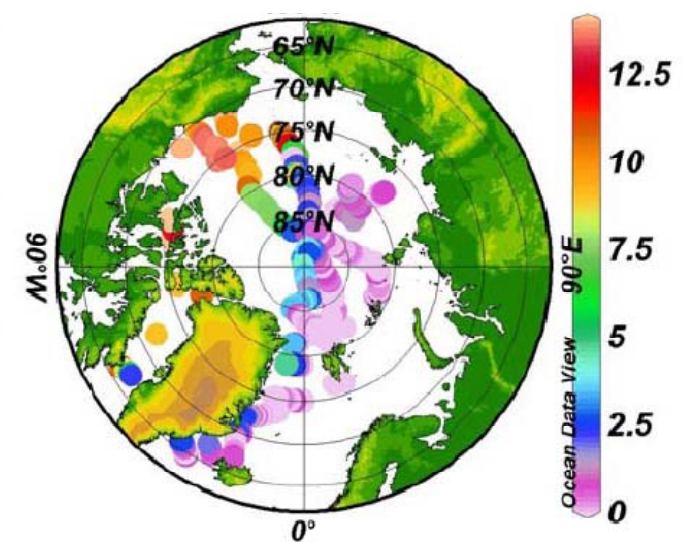
Sea ice melt is abundant in the Canada Basin and north of Svalbard.



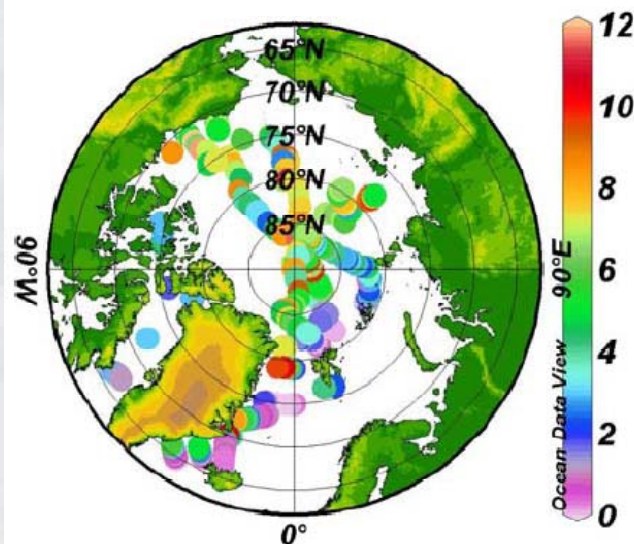
Total Freshwater



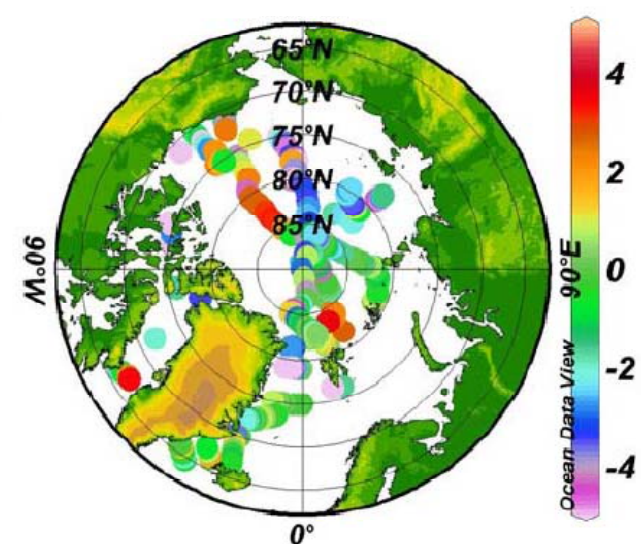
Pacific Freshwater



River Runoff

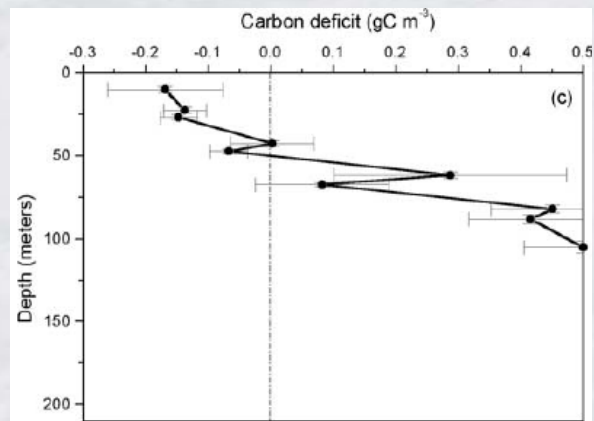


Sea Ice Melt

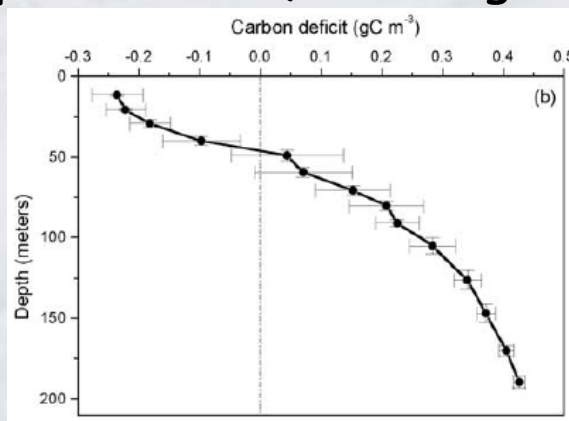


This information of surface water sources can be used to determine export production.

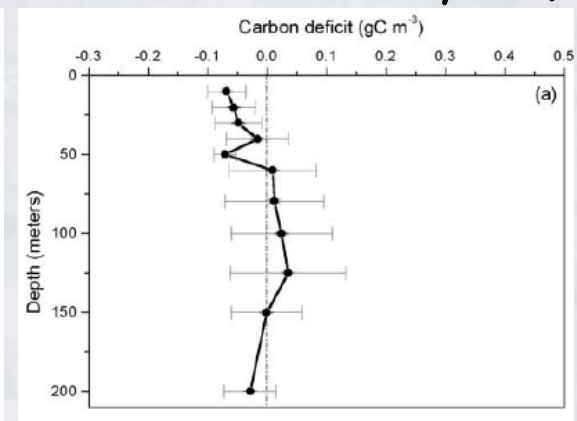
- Knowing the source waters contributing to a water makes it possible to determine the preformed concentration of e.g. phosphate.
- Comparing this to the measured concentration one can compute the deficit, i.e. what has been consumed.
- As the only sink of phosphate is primary production this deficit can be converted to new production (assuming residence time $\gg 1$ year).



Canada Basin
 5.8 gC m^{-2}



Makarov Basin
 8.4 gC m^{-2}



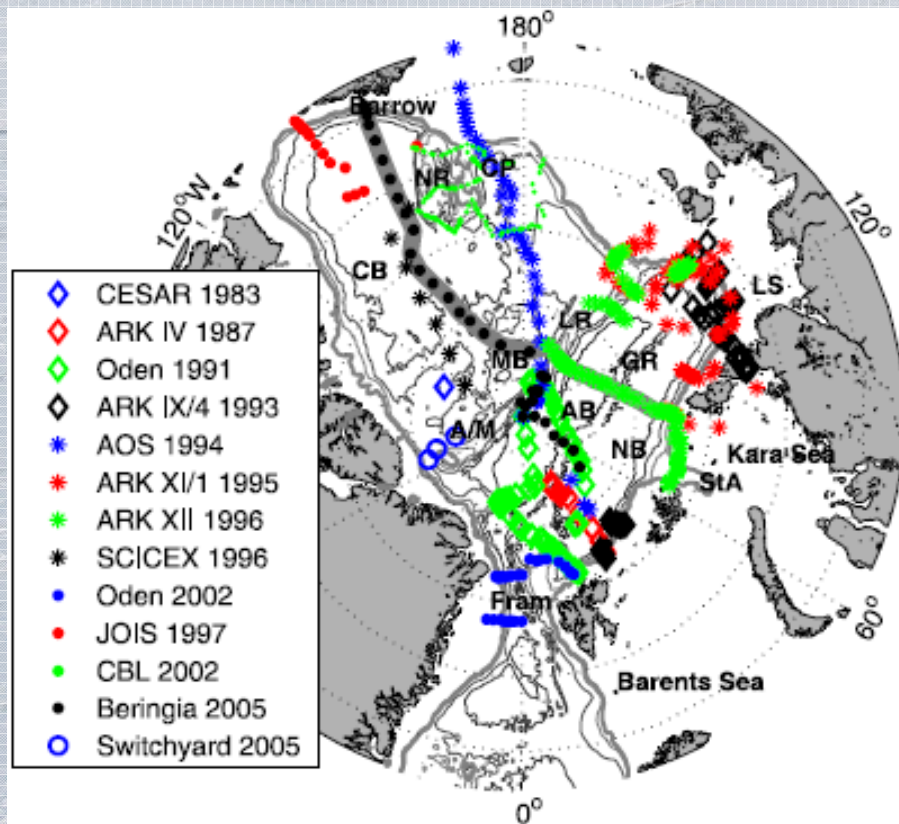
Eurasian Basin
 2.9 gC m^{-2}

With 15 years residence time this correspond to
 $\sim 0.5 \text{ gC m}^{-2}$ in the central Arctic Ocean

The anthropogenic carbon concentration by applying the transit time distribution (TTD) method.

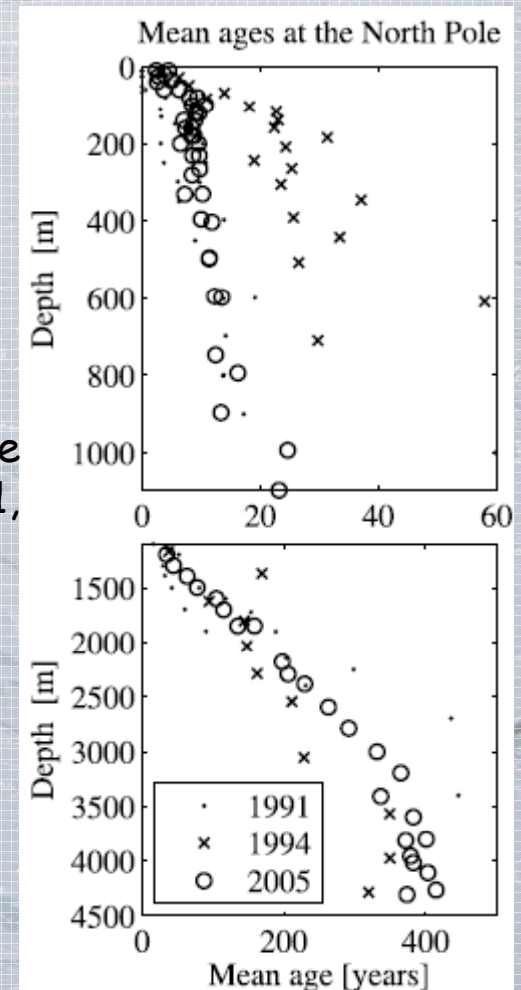
It is done by determining the concentration, c , of a passive tracer at any point, r , at any time, t , with knowledge of the TTD and the input function of the tracer at the sea-surface, with $G(r,t)$ being is the TTD.

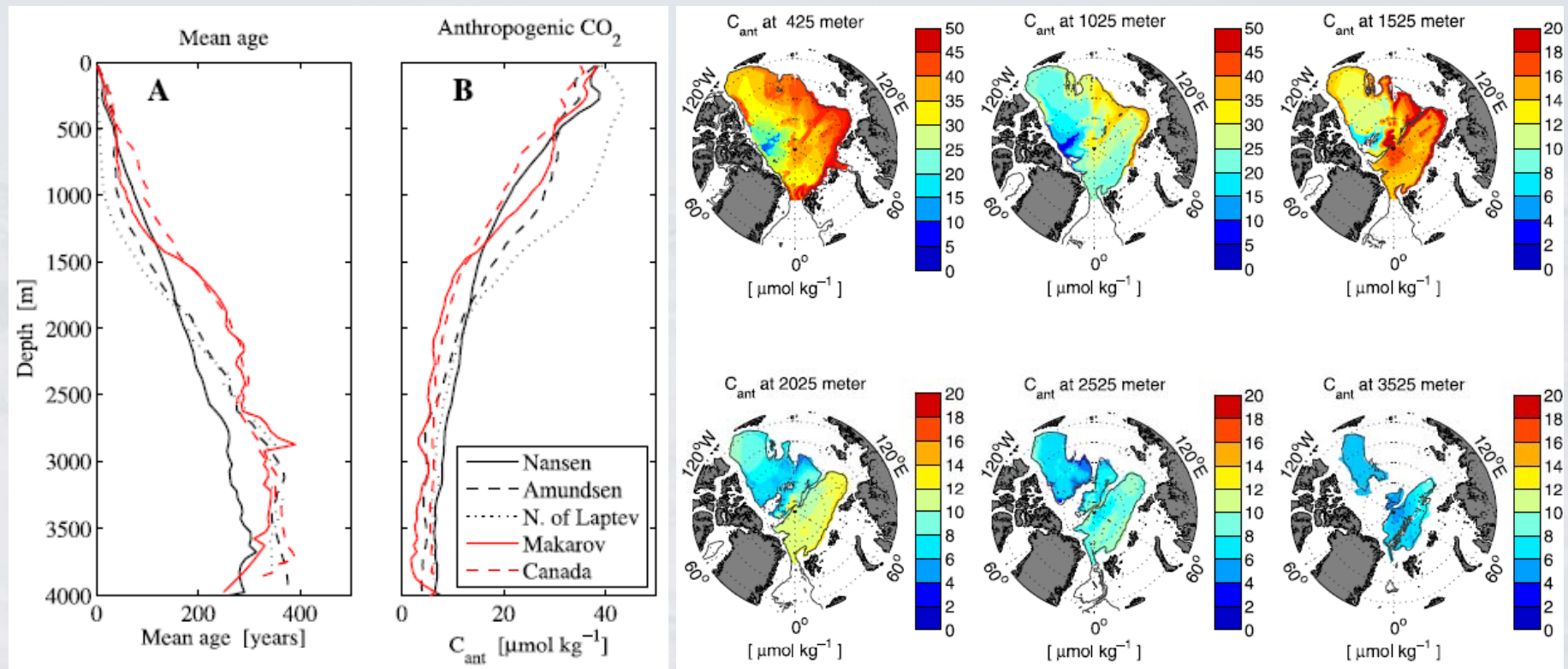
$$c(r,t) = \int_0^{\infty} c_0(t-t')G(r,t')dt'$$



Mean ages at the North Pole: 1991, 1994 and 2005.

Tanhua et al.
JGR, 2009

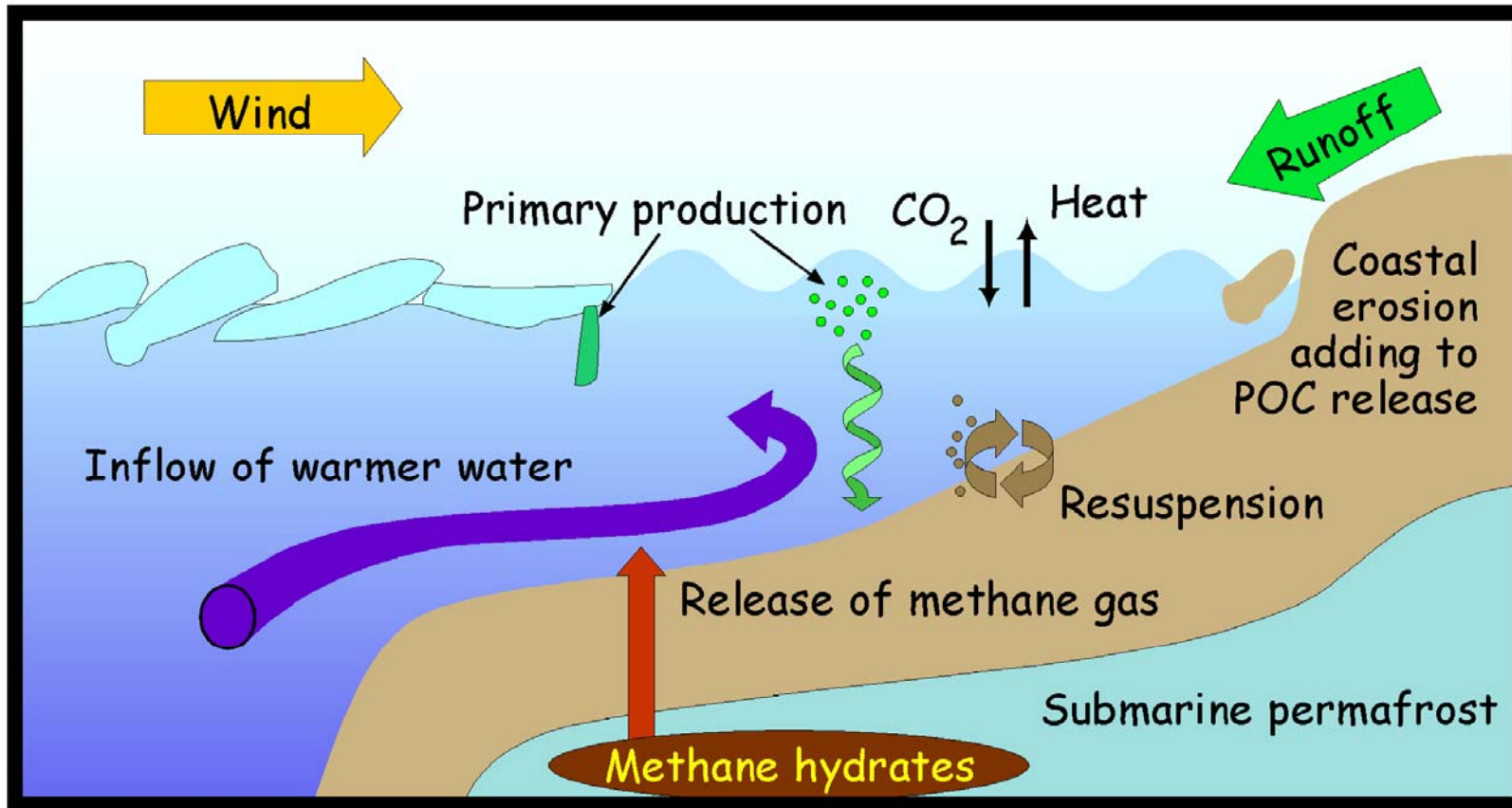




Applying the TDD method gives an inventory in the Arctic Ocean of $2.5 \pm 0.4 \times 10^{15}$ gC including the shelf seas, which is $\sim 2\%$ of all oceanic sinks.

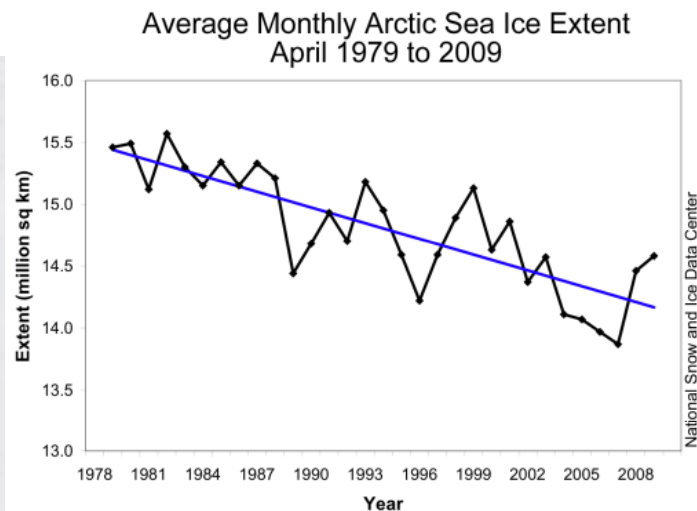
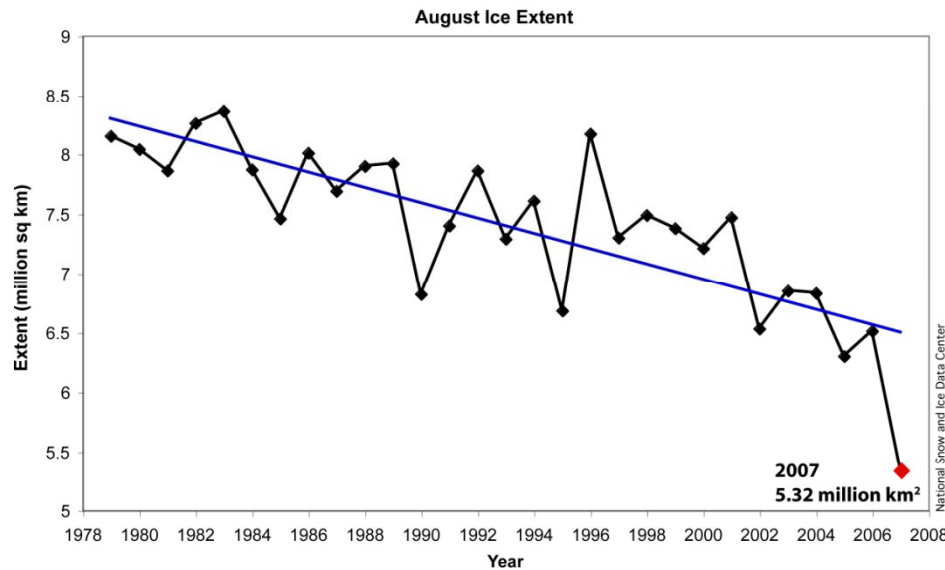
Tanhua et al.
JGR, 2009

Most C-transformation occurs over the shelves

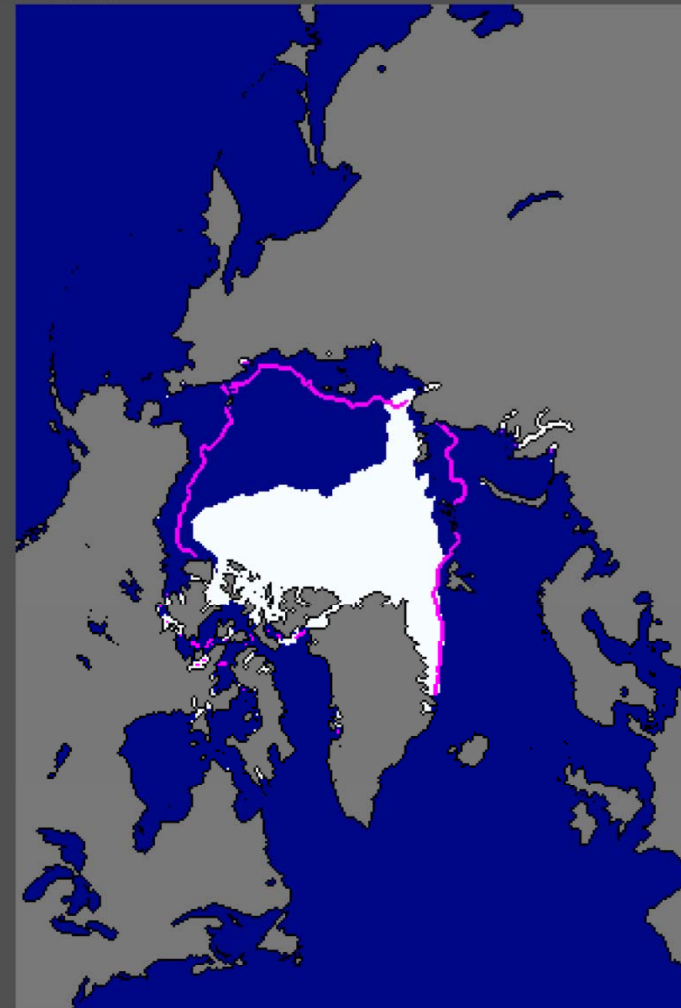


Includes biological production and decay, water mass modification, air-sea exchange, sedimentation and resuspension, terrigenous input, etc.

Will decreasing summer ice coverage impact the carbon transformation and fluxes?



Current Ice Extent
09/25/2007

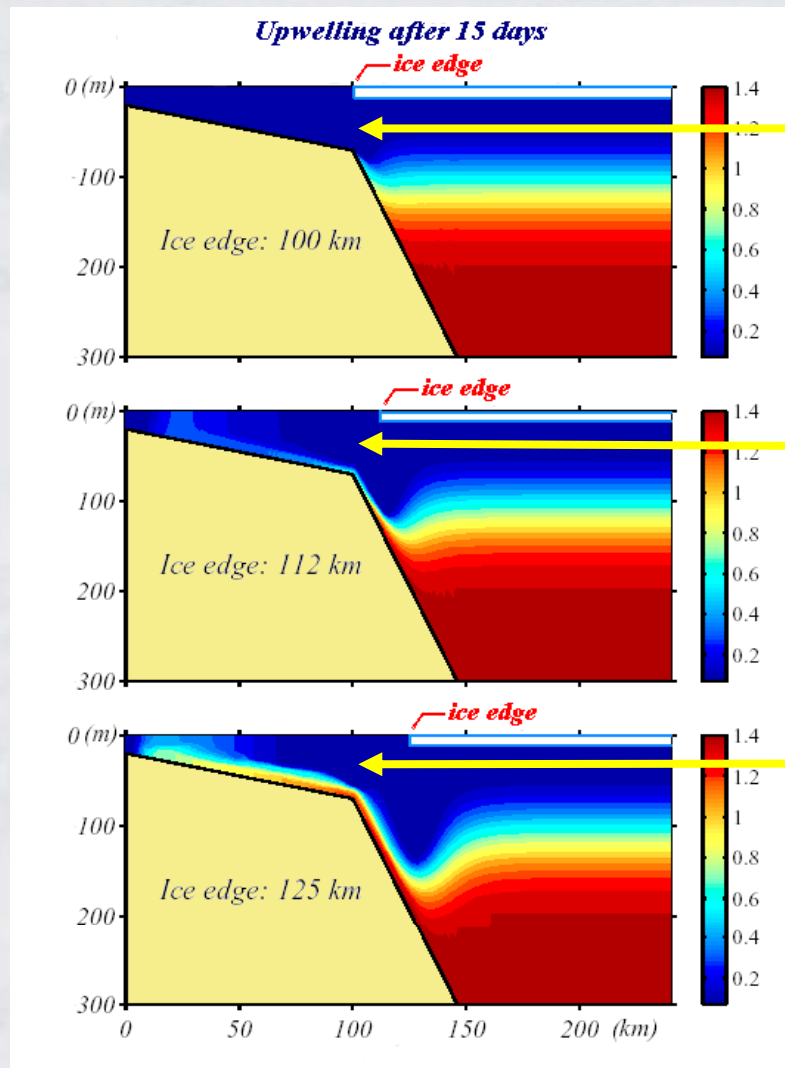


National Snow and Ice Data Center, Boulder, CO

median
ice edge

Total extent = 4.2 million sq km

Decreased ice cover will potentially increase upwelling, and thus also increase nutrient supply to the photic zone.



Model Results:

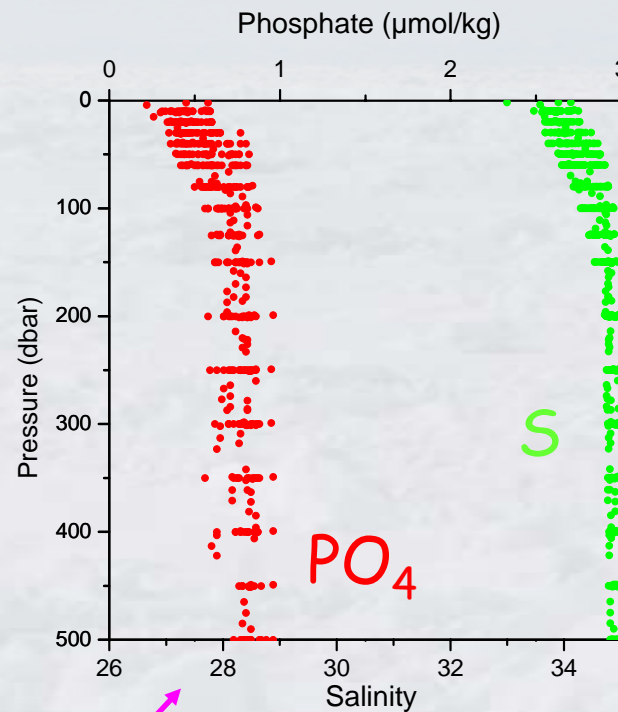
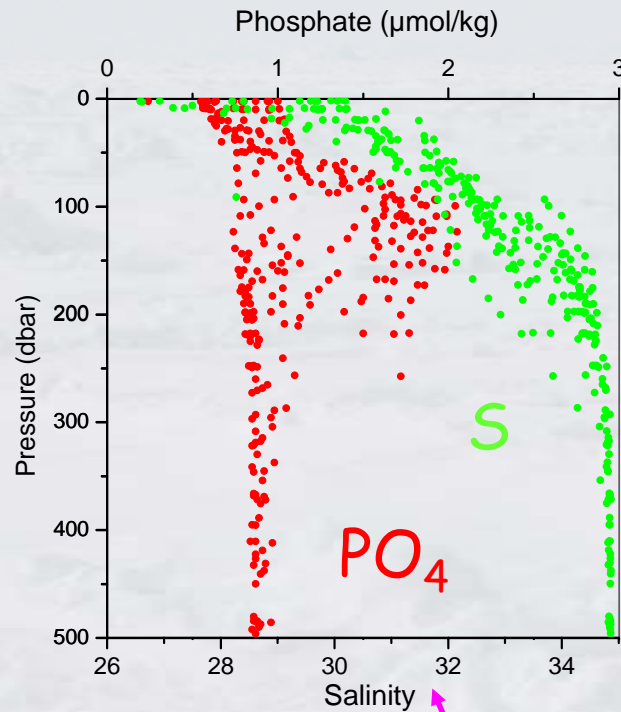
A circulation 'cell' forms with upwelling at the coast and downwelling at the ice edge. Only shelf water circulates if the shelf-break is ice-covered.

Continued retreat of summer ice cover exposes more and more of the shelf-break for longer periods of time to upwelling favorable winds.

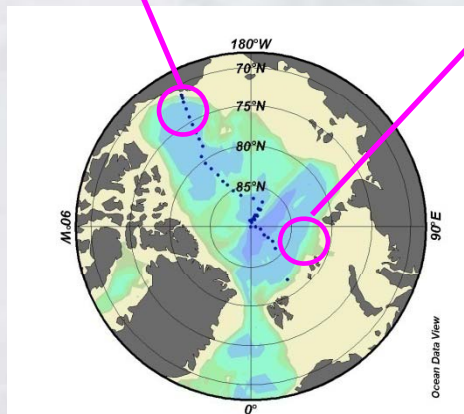
Upwelling depth increases as slope waters become ice-free. Salty, nutrient-rich water can now cross the whole shelf in a thin bottom boundary layer.

Carmack & Chapman, GRL, 2003

Decreased ice cover → increased upwelling → increased nutrient supply

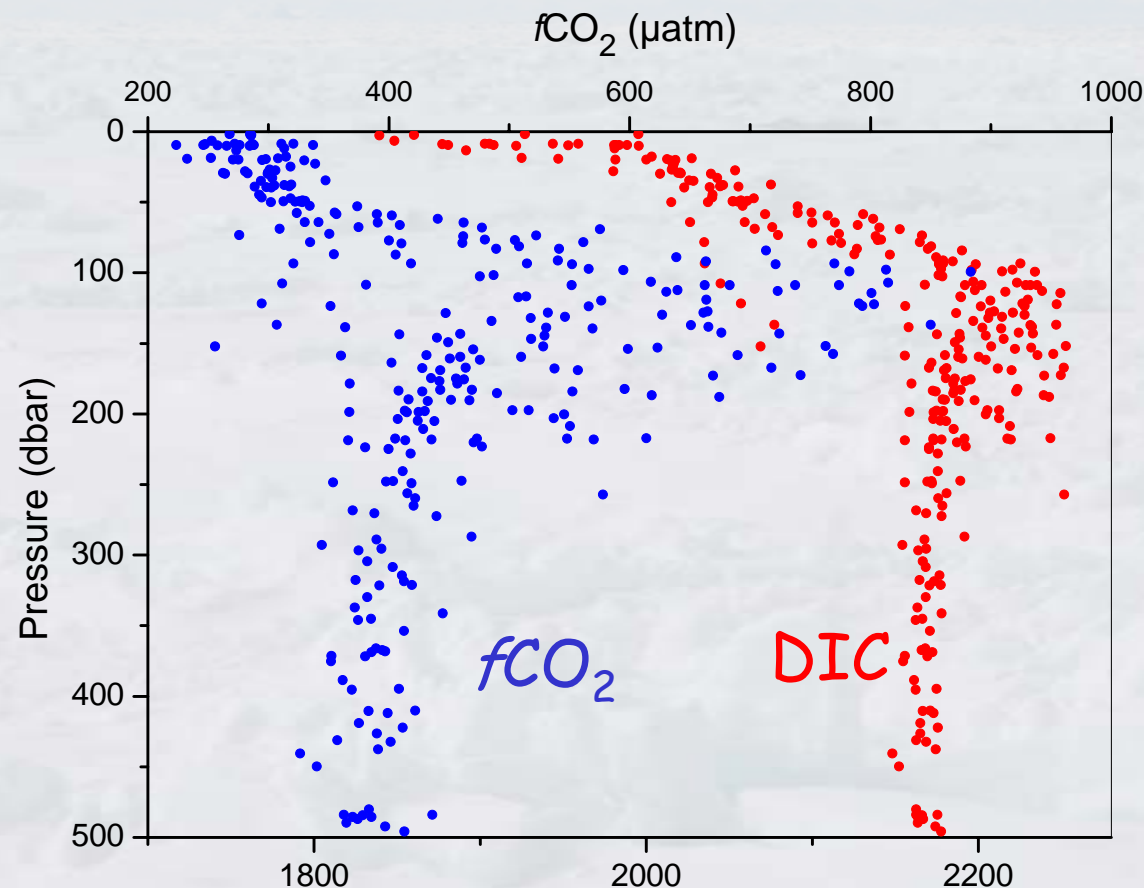


Pacific side;
nutrients higher
at ~100 m, but
profile more
stratified
compared to
Atlantic side.



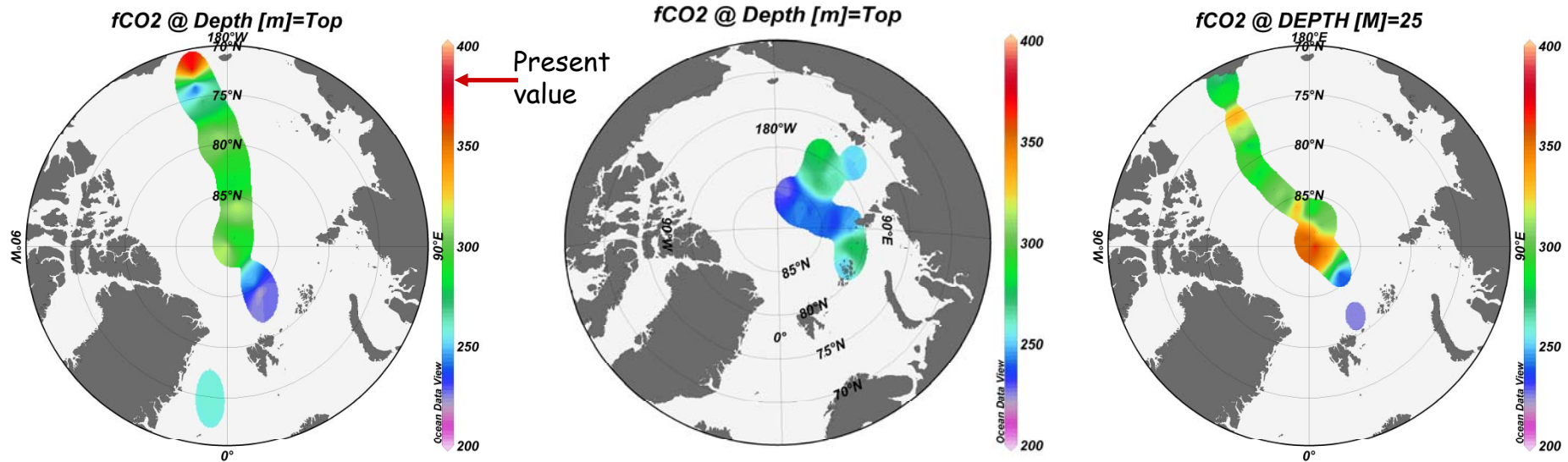
Beringia 2005 & Polarstern 1996

Decreased ice cover → increased upwelling → increased nutrient supply, but also → increased $p\text{CO}_2$ supply, which compensate the oceanic C-uptake!

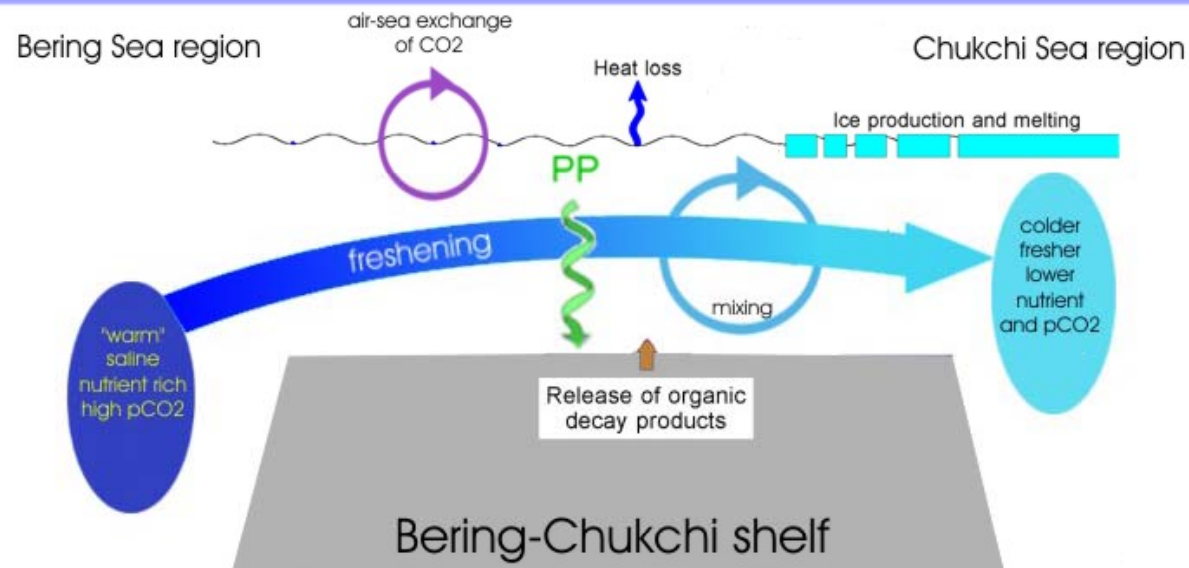


Also, if nutrients are mixed up at the Pacific side it will be followed by DIC, which compensate the effect of a potential increase in PP.

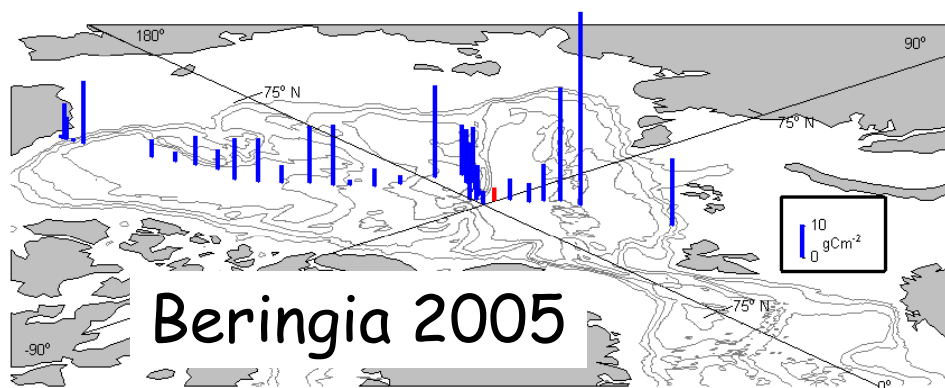
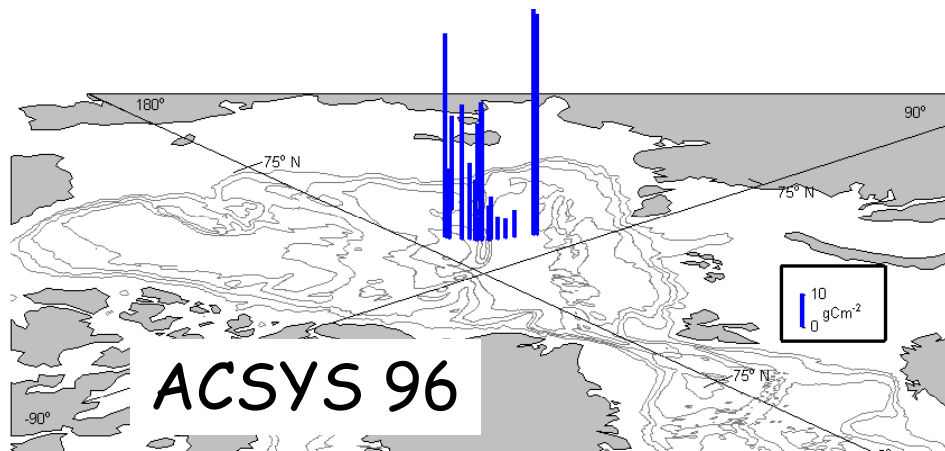
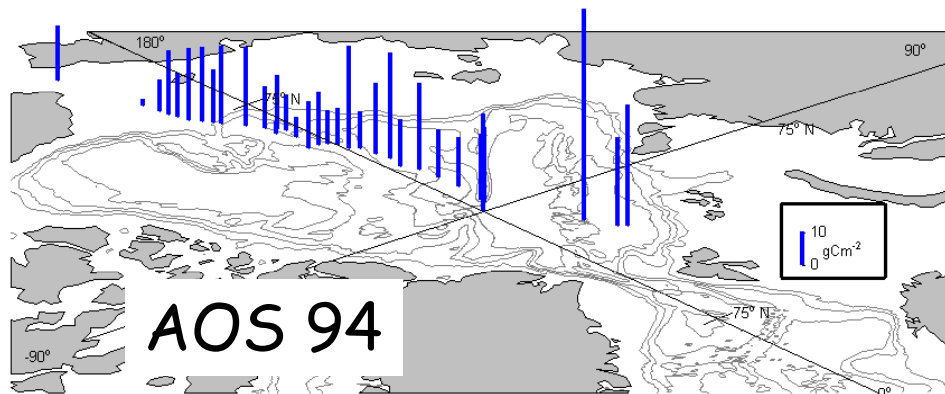
Beringia 2005 Dissolved Inorganic Carbon ($\mu\text{mol/kg}$)



The $p\text{CO}_2$ of the surface waters in the Arctic Ocean calculated with data collected during the cruises from left: AOS 94, ACSYS 96, & Beringia 2005.

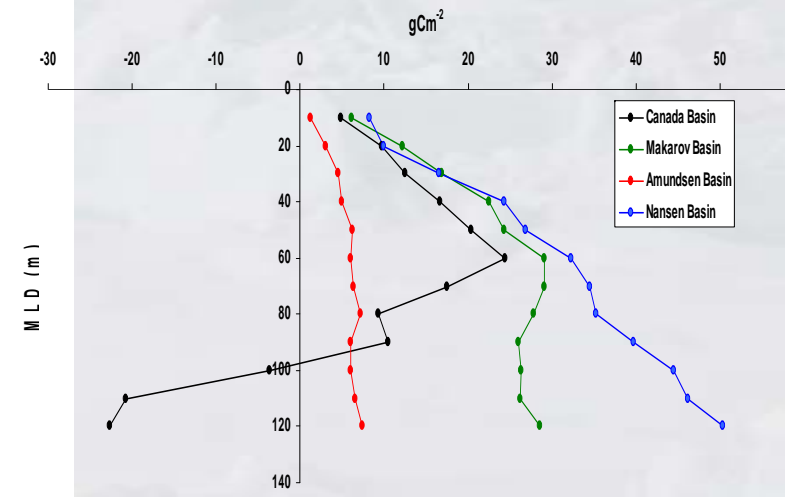


Or Barents Sea



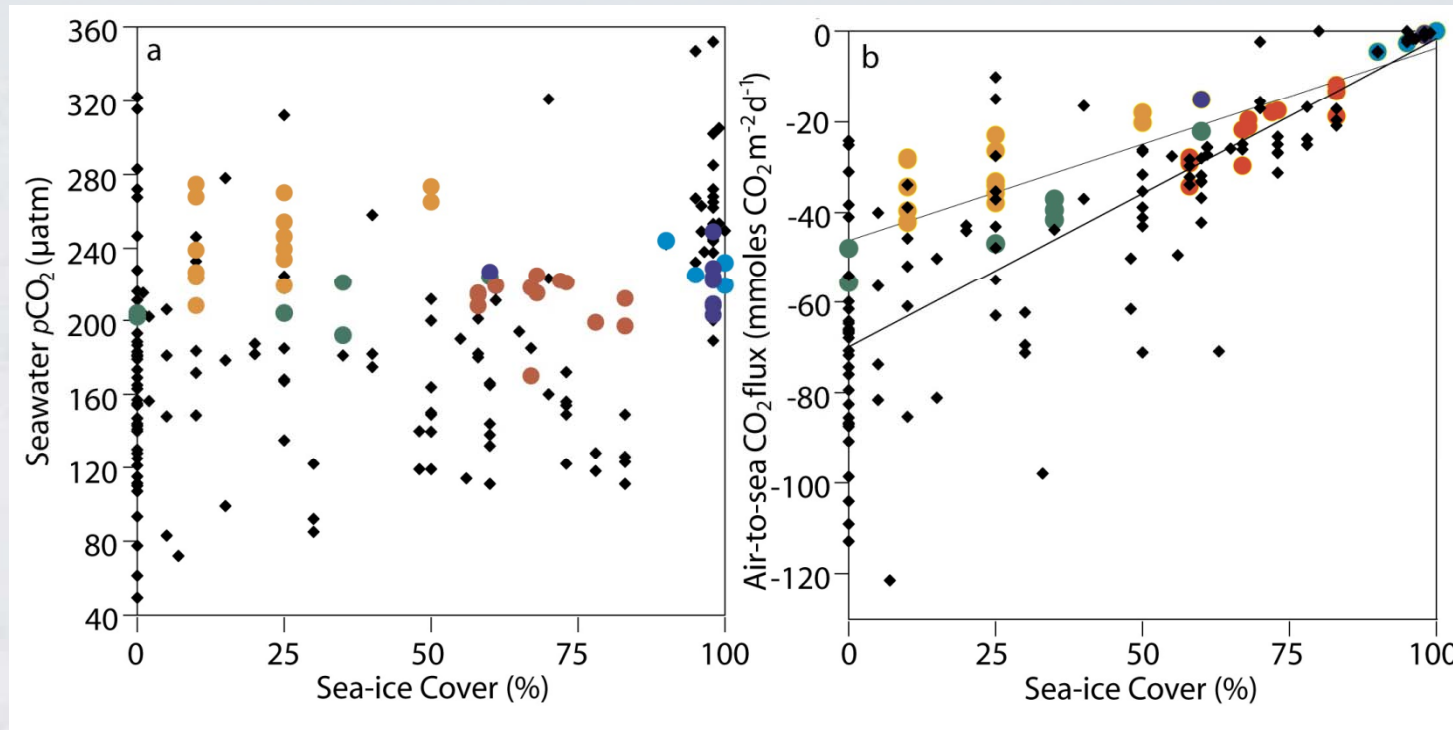
The potential uptake capacity of carbon in the winter mixed layer of the Arctic Ocean based on the pCO_2 under-saturation.

The potential uptake varies with time and location, but the average is close to 10 gC m^{-3} , which corresponds to about $50 \times 10^{12} \text{ gC}$ over the deep central basins.



pCO₂ in the Chukchi Sea region and the impact of Ice Cover on Air-Sea flux

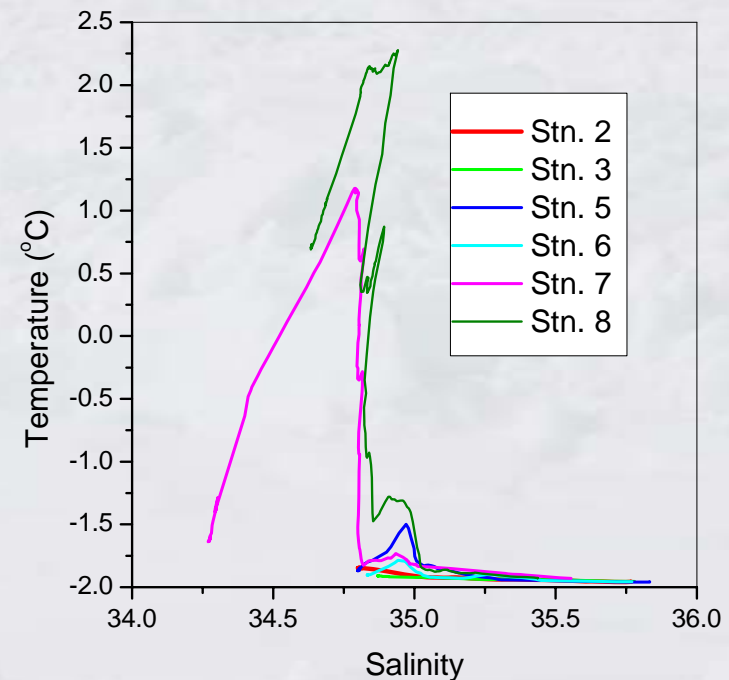
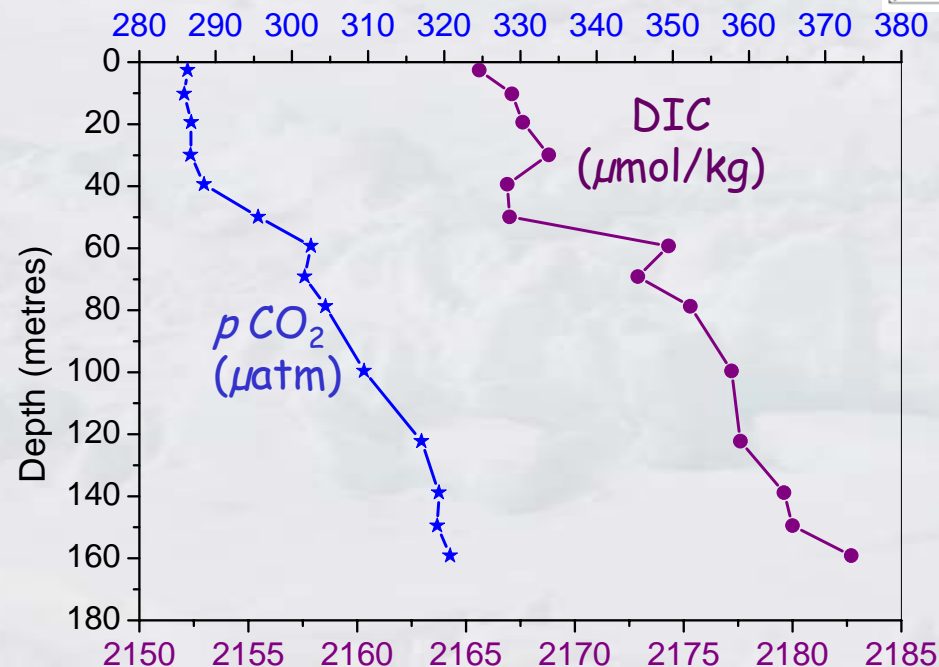
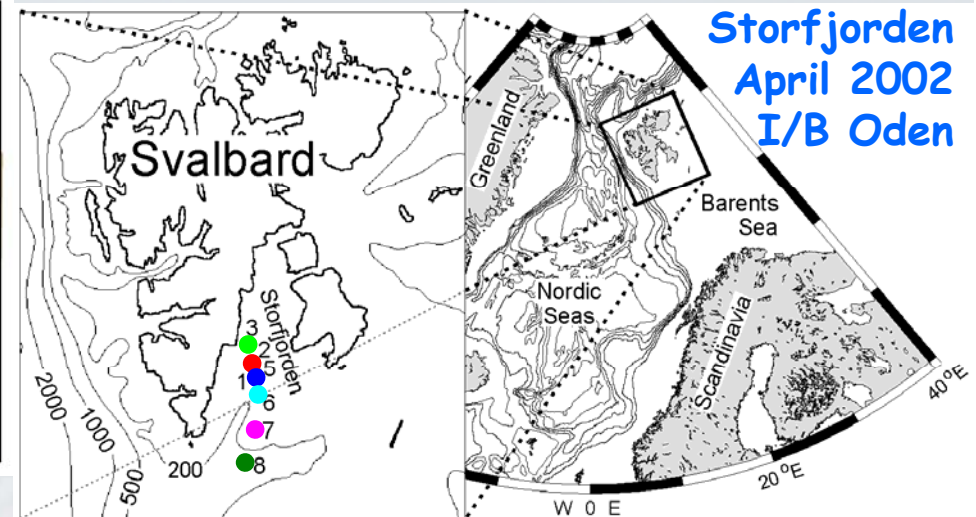
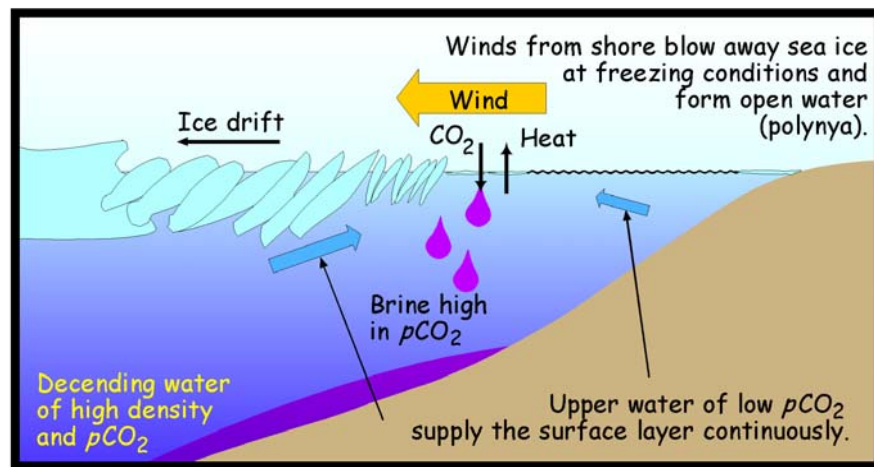
Bates et al. GRL 2006



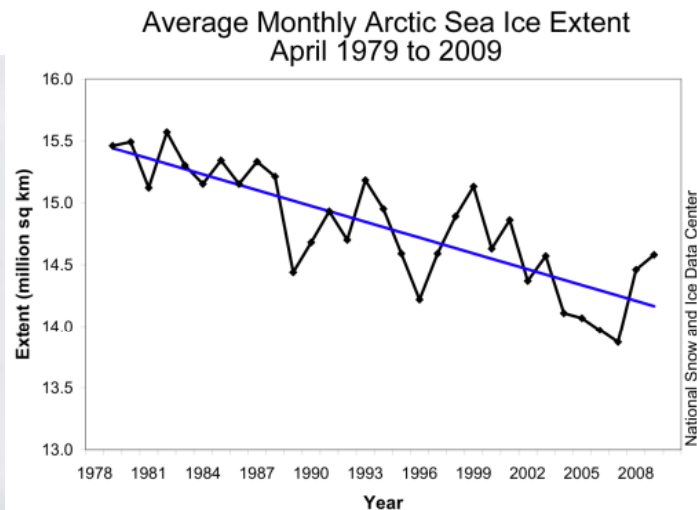
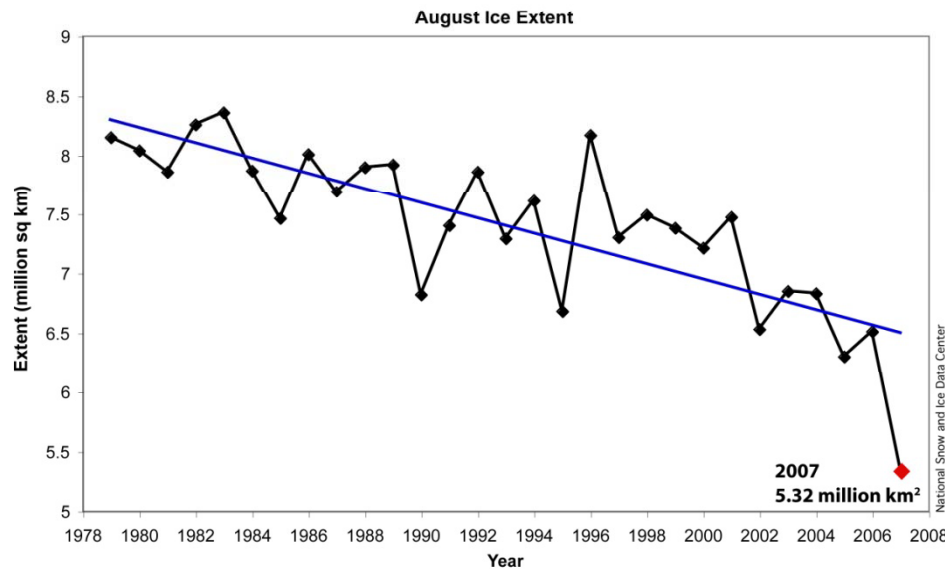
Large areas of under-saturated surface water, but no trend with ice cover. But as the flux much depends on the transfer energy at the sea surface it varies from zero in ice covered regions to ~60 mmol CO₂ m⁻²d⁻¹ (~0.7 gC m⁻²d⁻¹) in ice free waters.

Hence a decrease in ice cover has a major impact on the flux!

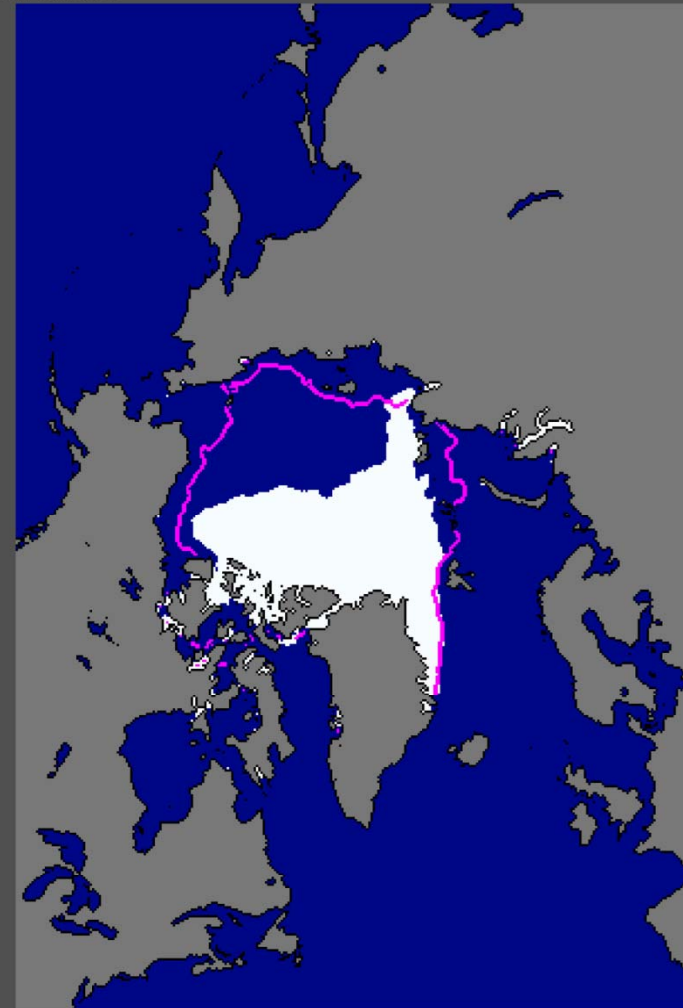
Sea Ice production, brine formation and gas exchange



Decreased summer ice coverage → larger ice production during freezing season → more brine production → increased CO_2 uptake → increased ventilation



Current Ice Extent
09/25/2007



National Snow and Ice Data Center, Boulder, CO

median
ice edge

Total extent = 4.2 million sq km

Arctic Ocean carbon fluxes and Climate Change

Effects on C-flux:

✓ more runoff \Rightarrow $\left. \begin{array}{l} \uparrow \text{stratification} \\ \uparrow \text{nutrient supply} \\ ? \text{light condition} \end{array} \right\} \Rightarrow \text{Primary Production?}$

✓ less sea ice \Rightarrow $\left. \begin{array}{l} \uparrow \text{light condition} \\ \uparrow \text{nutrient supply by increased vertical mixing} \end{array} \right\} \Rightarrow \text{PP?}$

✓ increased T \Rightarrow changed biology(?)

Future research:

- + There is a data deficit in the Arctic Ocean and hence all new data is welcome
- + Seasonal signal is substantial over the shelves and thus there is a need for seasonal data
- + This also goes for the surface of the central basins
- + Specific area with no (nearly) data is where the Lomonosov Ridge meet Greenland/Canada
- + Any specific data that is missing? Probably....