Warming and freshening of the Arctic Ocean in the 2000s

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Arctic Ocean takes part in the global circulation





Østerhus after Rudels

Strength of Atlantic meridional overturning and overflow density



Anomalies in the Subpolar and Subtropical North Atlantic



Hátún et al., 2005: Changes of the Subpolar Gyre





Curry et al., 2005 Changes of the Subtropical Gyre

4

Northward propagation of warm, saline anomalies



5 Holliday et al., 2008

1000

Strait

Sørkapp

Temperature in the Fram Strait Atlantic Water from summer CTD surveys



Increase of AW temperature and salinity, drop in density







7

10 years measurements of oceanic fluxes through the Fram Strait







Contribution to international ASOF (Arctic-Subarctic Ocean Fluxes)

Combination of international mooring programmes



• Time averaging, linear interpolation across section

Annual mean velocity, cross section component

Annual mean temperature

2002-2006: 17 moorings, 5 levels









Volume fluxes through Fram Strait – monthly mean values



- Northward/southward flow O(10 Sv)
- Barotropic
- Banded structure topographic influence

• ... Seasonal variability but no trend in volume transport

Computation of heat transport through Fram Strait only reasonable with the assumption of a stream tube



i.e. if water can be traced from inflow to outflow

• Heat transport through Fram Strait can be estimated only with assumptions ...



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Stream tube of Fram Strait inflow water through Arctic Ocean:

Inflow:

Only water with T>1°C considered.

Assumption: Can only return at same temperature or colder. Outflow:

Integrated until net volume flux is zero.

Atlantic Water (T>1°C) flow into the Arctic via the West Spitsbergen Current



Schauer et al., 2008 Schauer et al., 2009

Atlantic Water (T>1°C) flow into the Arctic via the West Spitsbergen Current



Additional 10 TW for 4 years = $12.6*10^{17}$ kJ

Schauer et al., 2008 Schauer et al., 2009

Variability of heat transport through Fram Strait

Computational:

12.6*10¹⁷ kJ corresponds to 4 m of ice melt in Eurasian Basin (10¹² m²)



S(t)

Heat transport through Fram Strait = $\frac{\partial H}{\partial t}$

Arctic warming

Quadfasel et al. 1991 Morison et al., 2000 Schauer et al., 2002 Morison et al., 2005:

. . .

17

Heat loss north of Svalbard?



Cokelet et al., 2008: October 2001 <u>520 W/m</u>² in upper 500 m over 500 km *100 km.

→ <u>26 TW</u>

Heat flux through Fram Strait before 2002 about <u>30 TW</u>.

Compares well, but 500 W/m2 considered too high by atmospheric community.

Heat loss in northern Fram Strait?



-20 -40-60-80-100-120-140-160-180-200 -220-240-260-280Su.

20

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Fieg et al., submitted

• What balances the variable heat transport through Fram Strait?

Heat flux through Fram Strait Before 2002 about 30 TW, after 2002 about 40 TW

=> Additional 10 TW for 4 years = 12.6*10¹⁷ kJ

$$S_{normal} = \oint_{(t)}^{Z(t)} dy \int_{0}^{Z(t)} c_{p} \cdot \rho \cdot v_{\perp}(y, z, t) T(y, z, t) dz$$

S: Sink

Changes of ocean properties in the Eurasian Basin since 1991





Assumption: Sections represent Basin Properties

Photo: Stefan Kern



Atlantic Water (warm core, 150 – 400 m)





AW boundary current warmer and more saline
Return flow in Nansen B warmer by 0.5K and more saline
Return flow in Amundsen B warmer than early and colder than late 1990s, warmer than 2005

Nansen Basin Area ca. 0.5 * 10⁶ km²



Heat increase = $2.5 \times 10^{17} \text{ kJ}$

Atlantic Water (lower part, 400 -1000 m)



Same tendencies in BC and Nansen Basin (dT=0.2K)
Similar T change in Amundsen B return flow



Heat increase = $2.4 \times 10^{17} \text{ kJ}$

Inflow of Barents Sea branch of Atlantic Water



Deep Water (1000 – 2500 m)



Increase by 0.02K in Nansen Basin



Decrease in salinity (approx. 0.01 psu) in both basins

Heat increase = $0.6 \times 10^{17} \text{ kJ}$



Bottom Water (2500 m to bottom (up to 4500 m))



Both basins warmer by 0.02 °C



Heat increase = $2*10^{17}$ kJ



Heat balance

Fram Strait additional 10 TW for 4 years = $12.6*10^{17} \text{ kJ}$

+S(t) = heat transport through Fram Strait

AW core 2.5E+17 kJ AW deep 2.4E+17 kJ EBDW 6E+16 kJ EBBW 2E+17 kJ **T.5E+17 kJ**

 ∂H

 ∂t

Increased heat loss to ice and/or atmosphere?

Refers to 60% of the additional heat transport Missing: Estimate of warming in the Canadian Basin.





Deep water warming through variable geothermal heating?



Upper fresh water layer (Polar water)



30

NAOSIM - Simulated Arctic fresh water content 1950 - 2000



Freshwater circulation

Shipborne, airborne, submarine CTD

IPY 2007/08: Deployment of 17 Ice-tethered platforms International effort





Freshwater accumulation in the Canada Basin



Increase of up to 11 m FW (Sref =34.8) as compared to EWG climatology

... causes strong surfcae geostrophic flow towards Fram Strait and Canadian Archipelago

McPhee et al., 2009

<u>Summary</u>

- Temperature anomalies advected from subpolar Atlantic
- Heat transport computation through Fram Strait possible only with fierce assumptions

=> 20 - 30 TW late nineties, increase by 30% during early 2000s

• Eurasian Basin warming explains 60% of the *increased* heat transport in the early 2000s (more if Canadian Basin warming is included)

• The remaining 40% (or less if ...) probably increase of the heat loss north and west of Svalbard (from 30 TW to 35 TW).

- Strong warming of deep and bottom waters: Other heat sources?
- Freshening of the Arctic during 2000s mostly in the central part