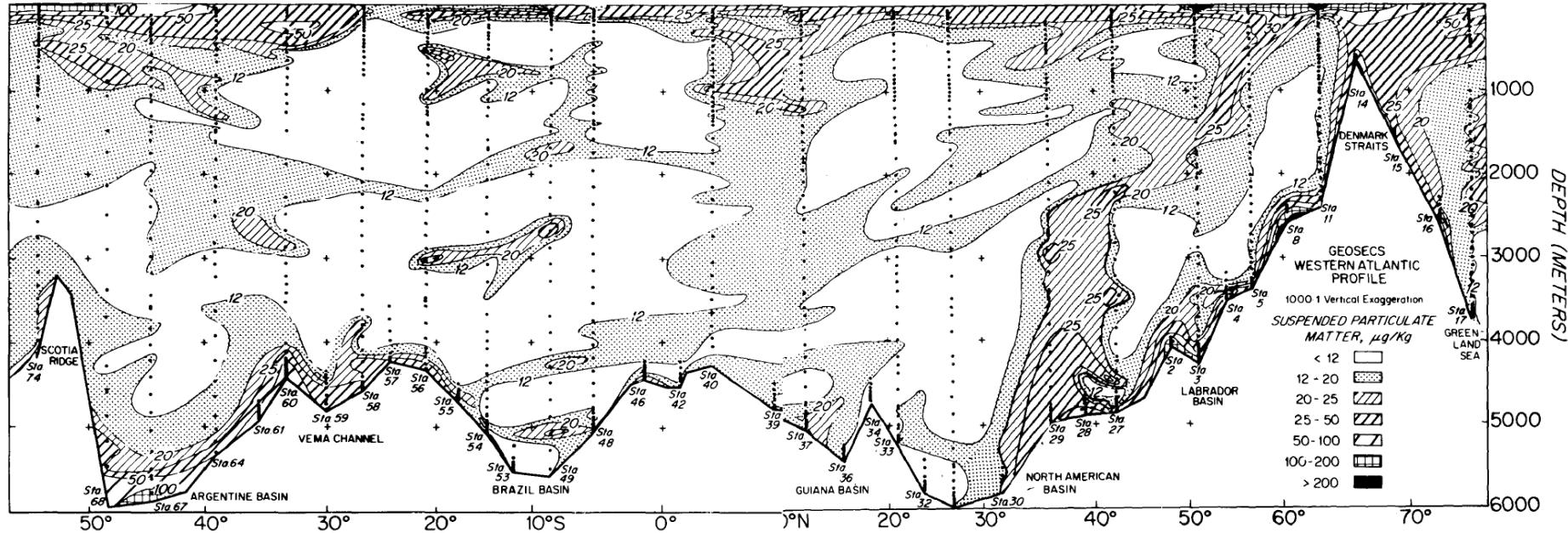


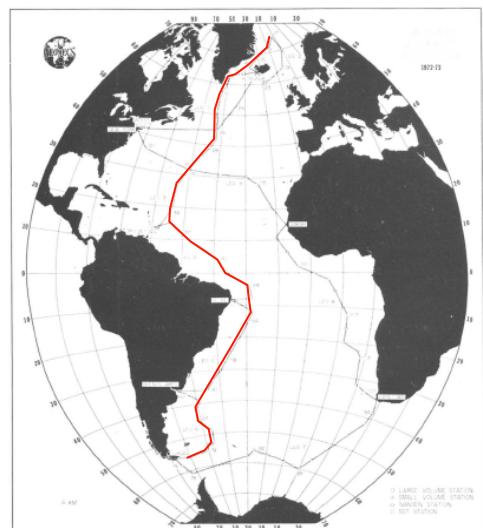
Bottom and Intermediate nepheloid layers

Michiel Rutgers van der Loeff and Sven Kretschmer

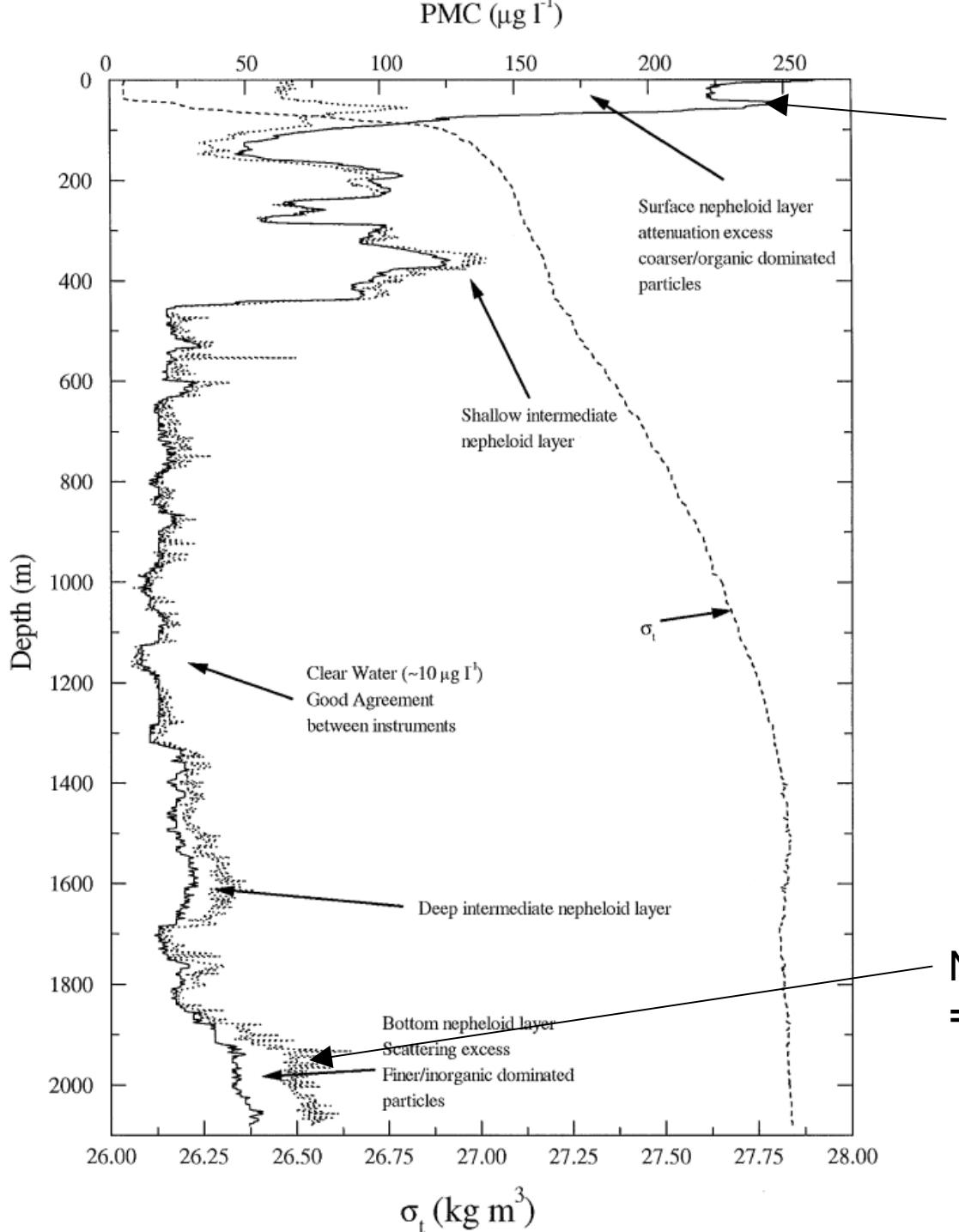
Suspended particulate matter during GEOSECS



Longitudinal section of the dry weight of particulate matter in the western Atlantic Ocean.



Brewer et al., 1976

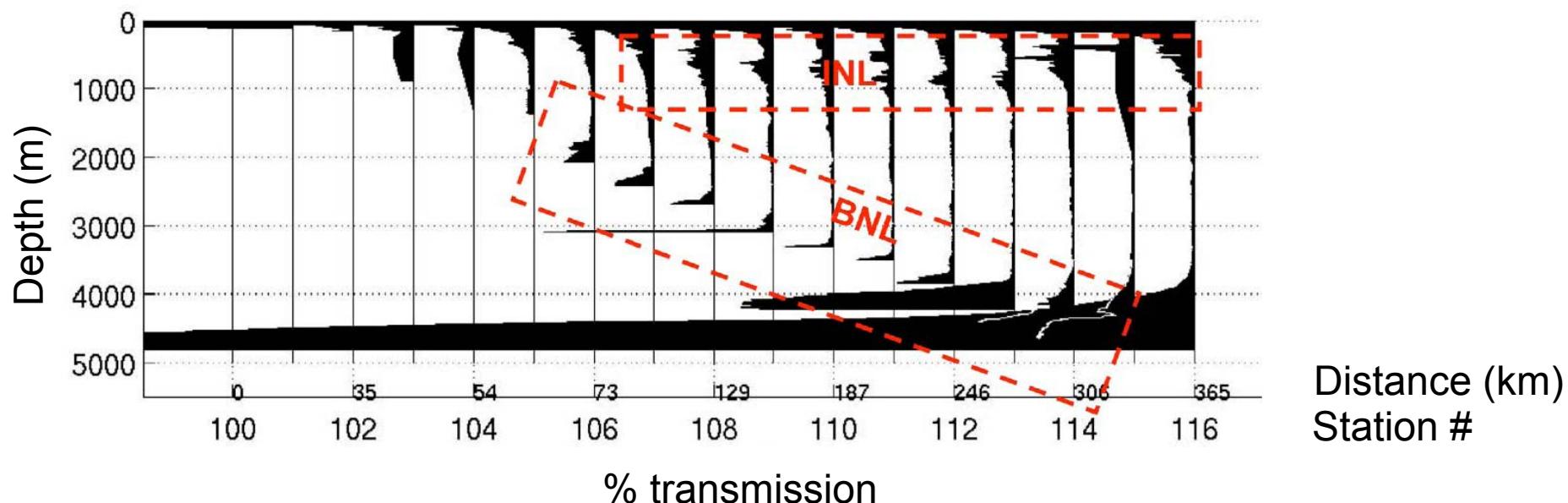


Transmissometer
= beam attenuation

Nephelometer
= scattering

Woods Hole

Bermuda

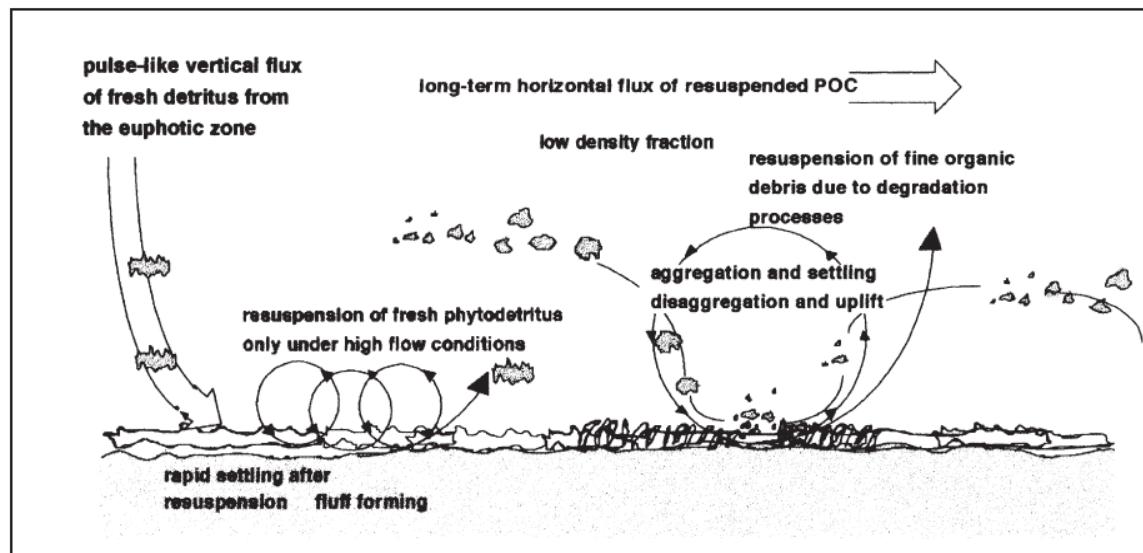


BNL and INL

- Bottom Nepheloid layers:
 - Within a basin (local resuspension)
 - Long-range transport
 - Tracer transport and exchange
- Intermediate Nepheloid Layers

Local resuspension

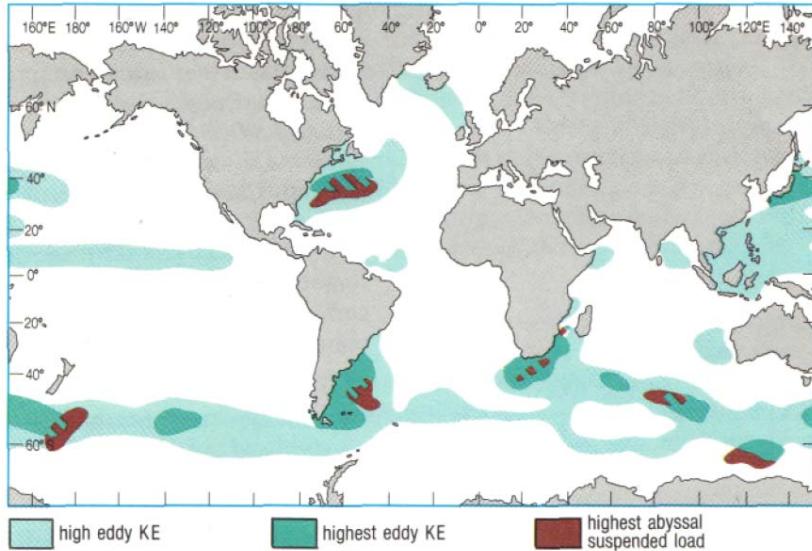
- Generation, distribution, size spectra in BNL described by McCave (1984, 1986, 2001)
- Vertical mixing in bottom layer studied with ^{222}Rn . Vertical extent enhanced by detachment of bottom mixed layers



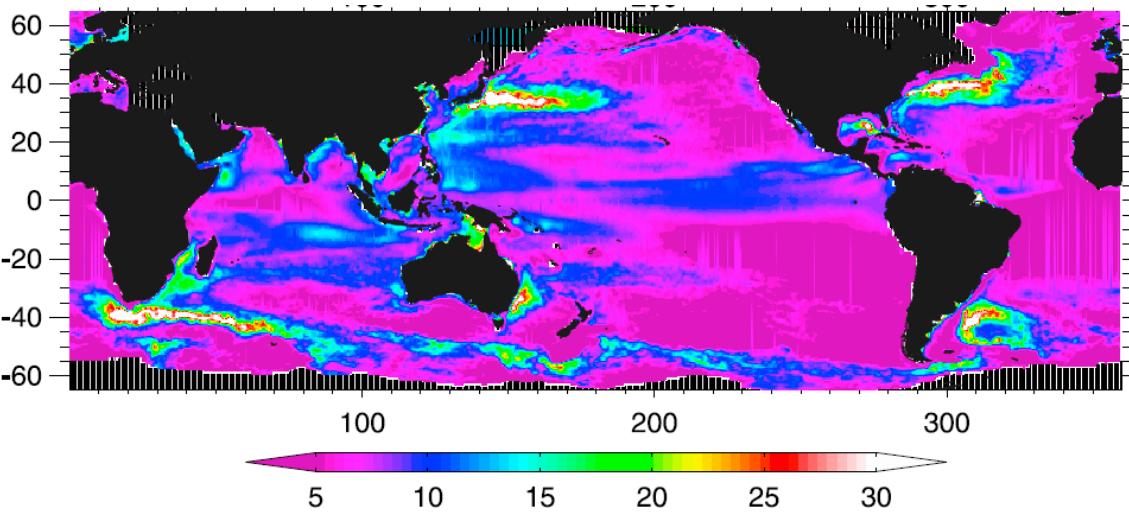
Settling and resuspension

Phytodetritus pulses
Aggregation/Disaggregation
Faunal activities

Resuspension by high eddy kinetic energy



Open University based on
Hollister and McCave 1984

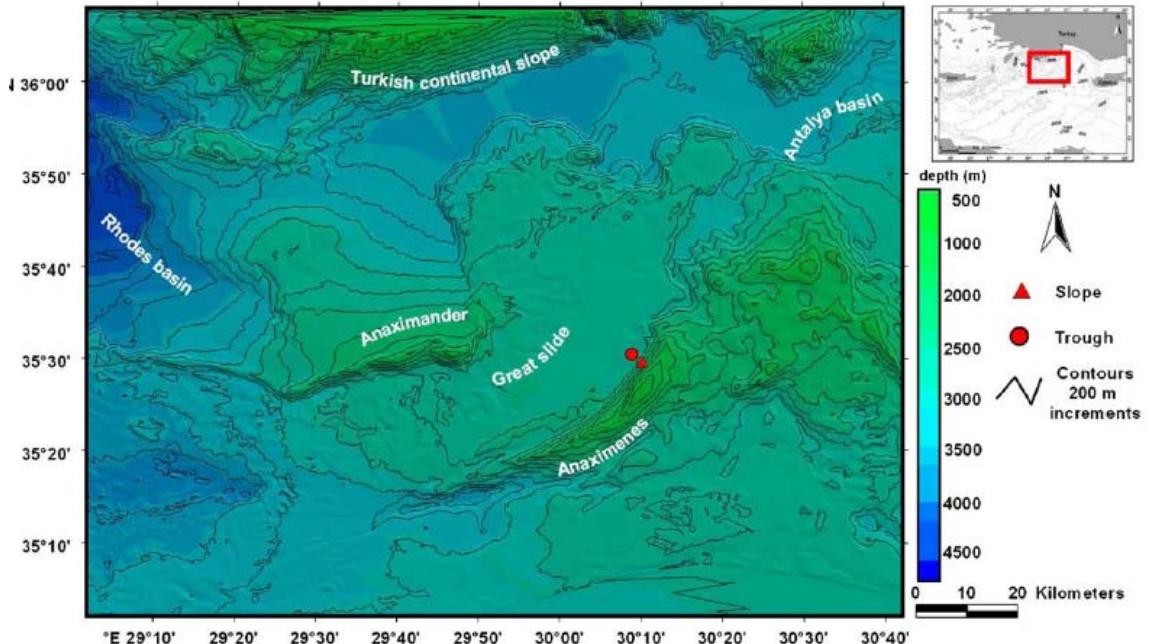
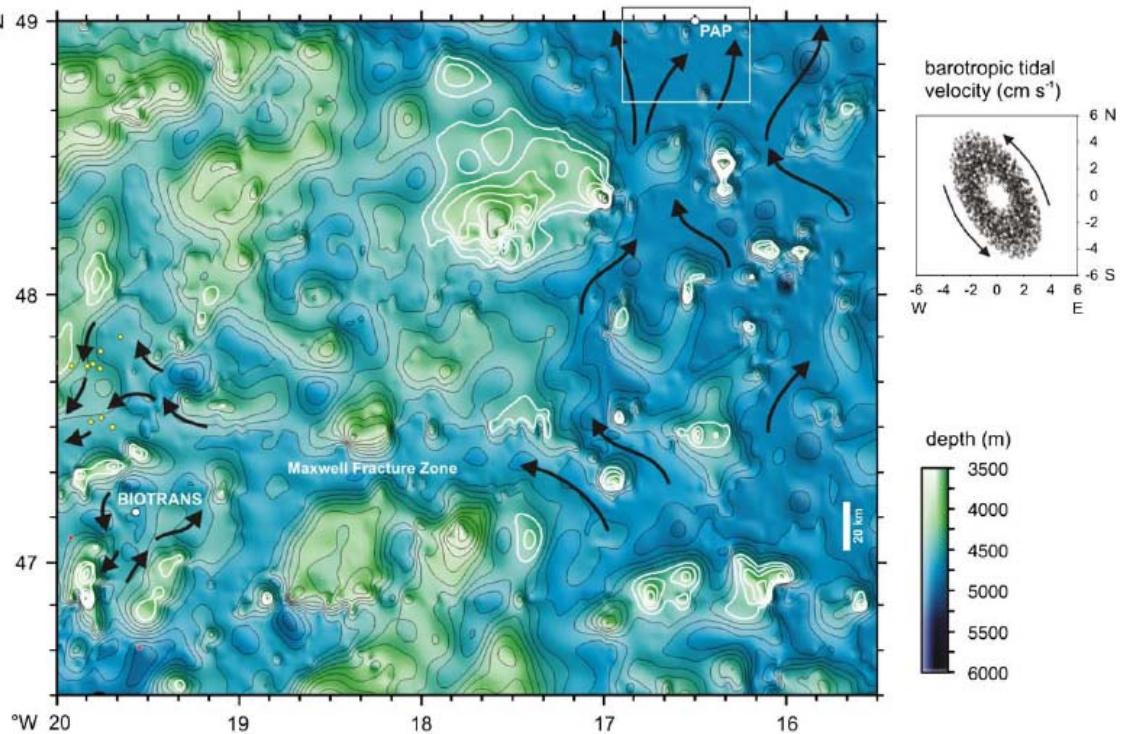


Standard deviation of sea
surface height (in cm) from
satellite altimeter data

Lee-Lueng Fu, JGR, 2009

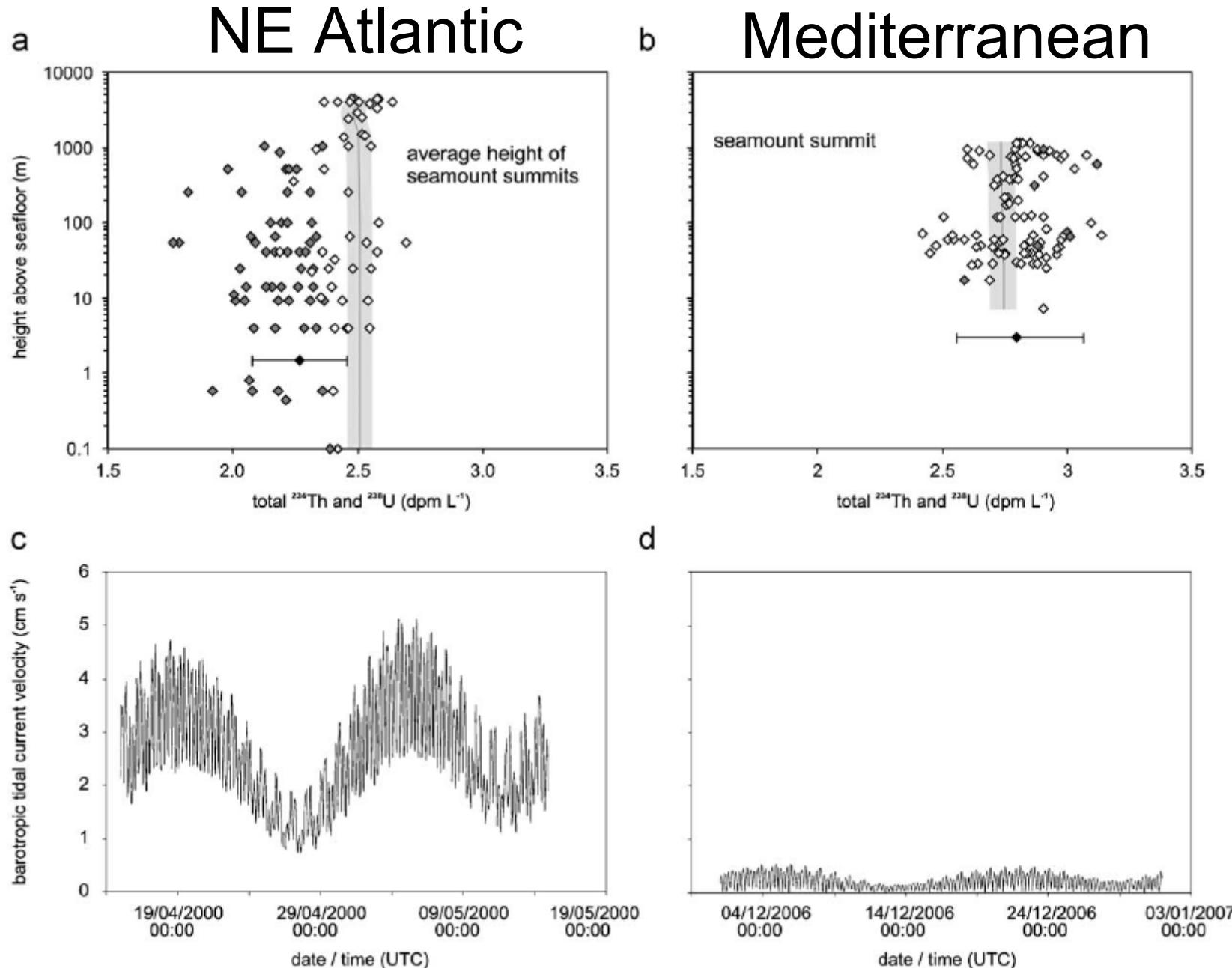
Resuspension by tides:

NE Atlantic

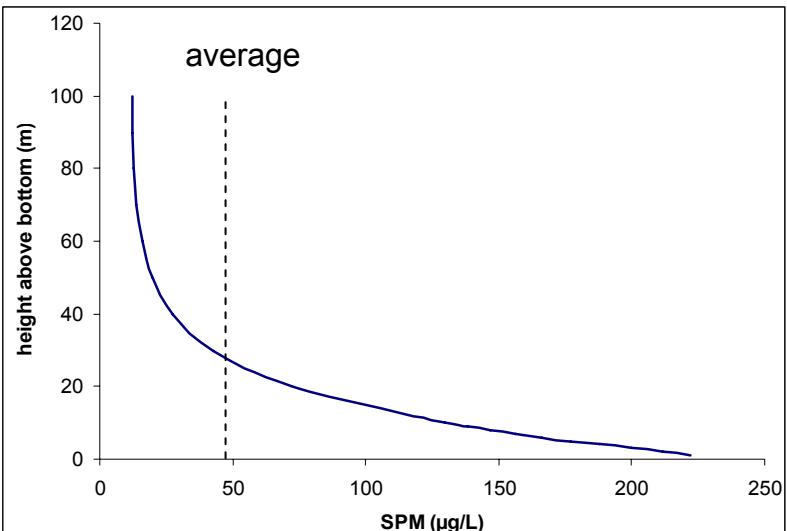
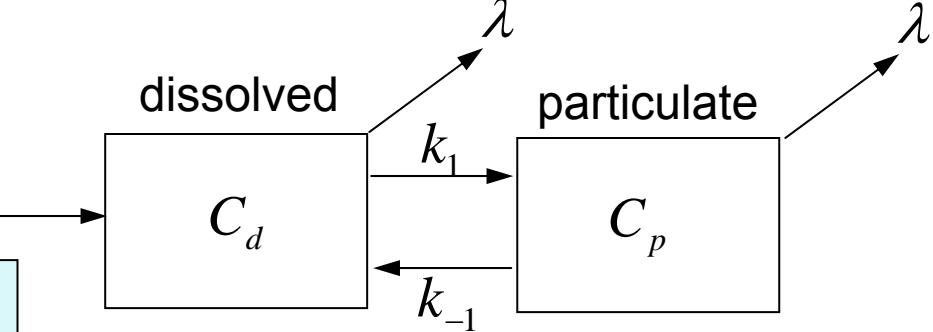
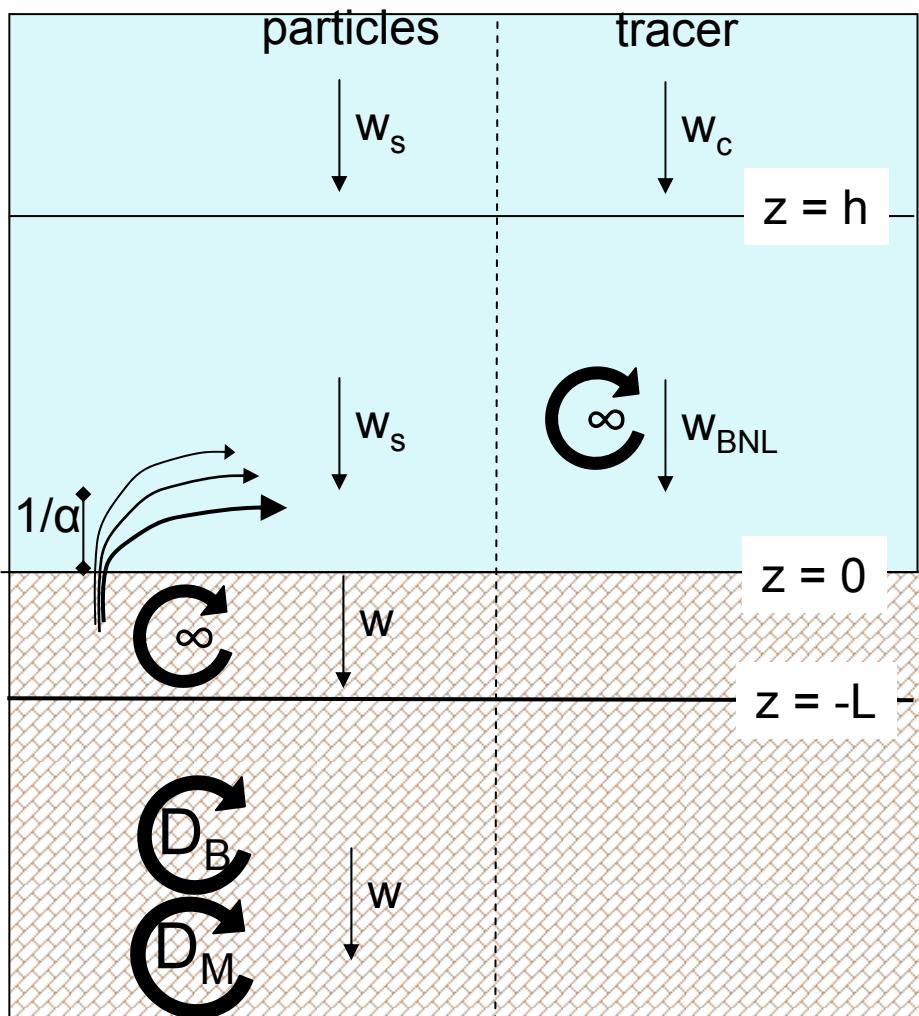


E. Mediterranean

Turnewitsch et al., 2008
Peine et al., 2009



Resuspension model

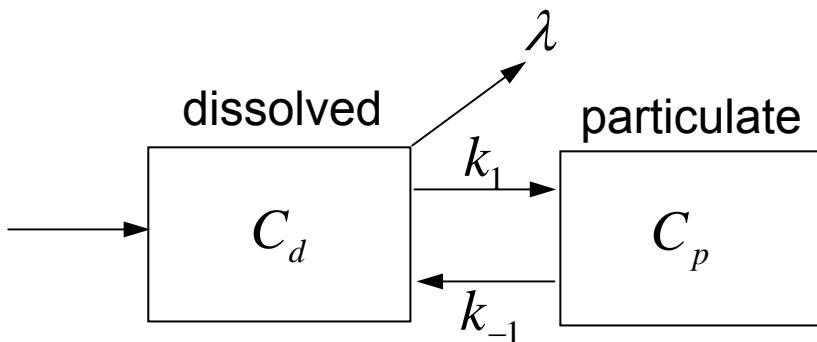


Apply to	Halflife (y)	Total activity BNL/clear water
^{234}Th	.07	0.86
^{228}Th	1.9	0.26
^{210}Pb	22.3	0.27
^{230}Th	75400	1.45

Bacon and RvdL, 1989; Boudreau, 1997; RvdL and Boudreau, 1997

Particle concentration effect

or: is K_d a constant?



Adsorption equilibrium exists when

$$\frac{C_p}{C_d} = \frac{k_1}{k_{-1} + \lambda}$$

$$\frac{X}{C_d} = \frac{k_1}{P} \frac{1}{k_{-1} + \lambda} = K_d$$

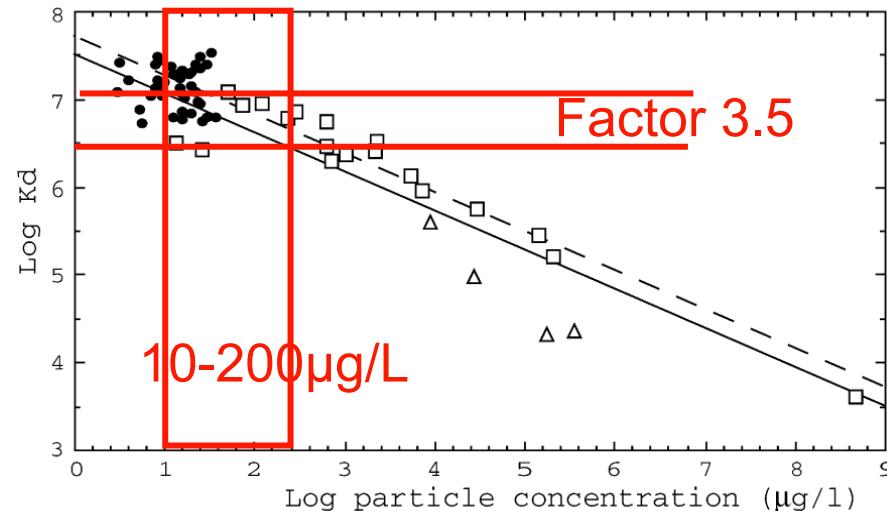
P = suspended load

X = activity per mass of particles

C_p = X.P =particulate activity

C_d = dissolved activity per mass of water

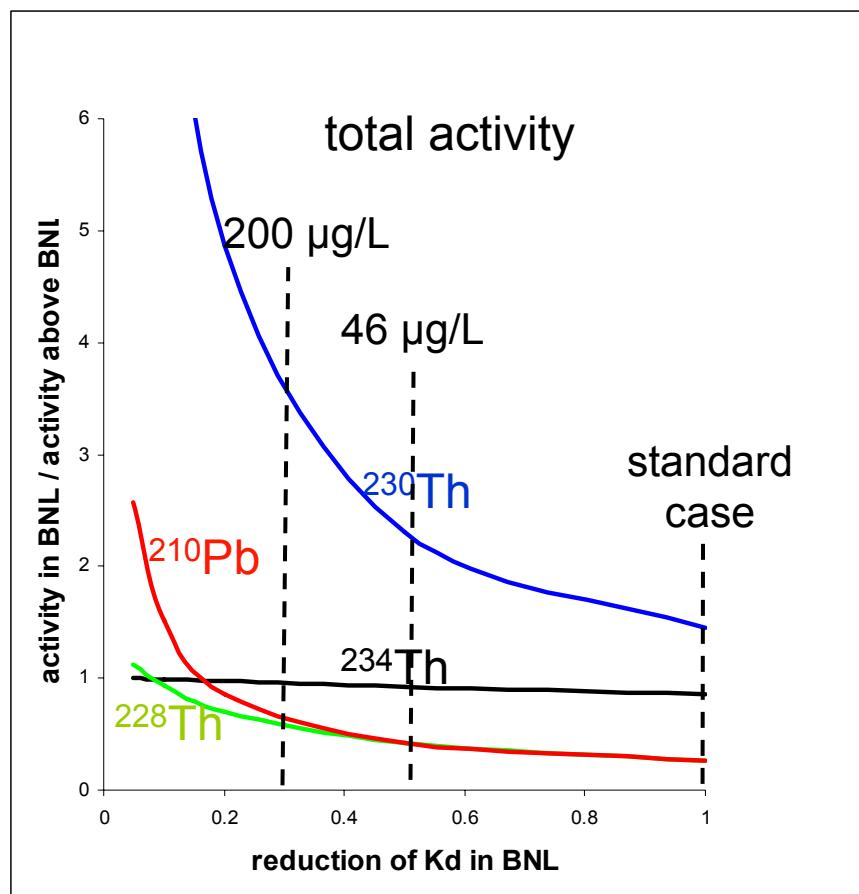
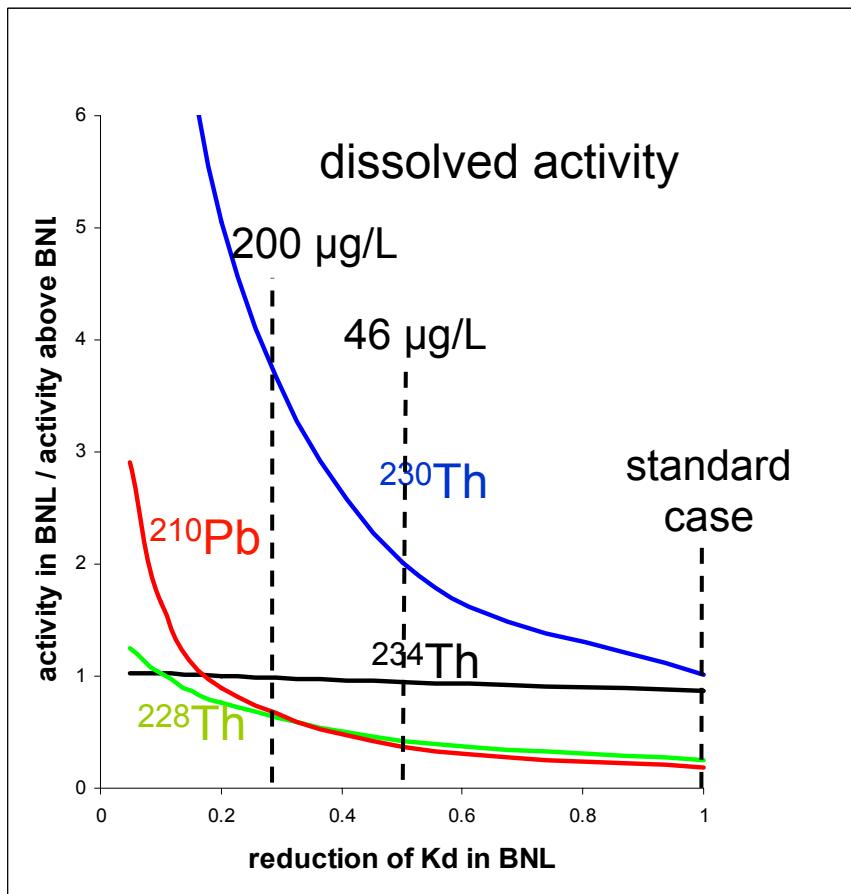
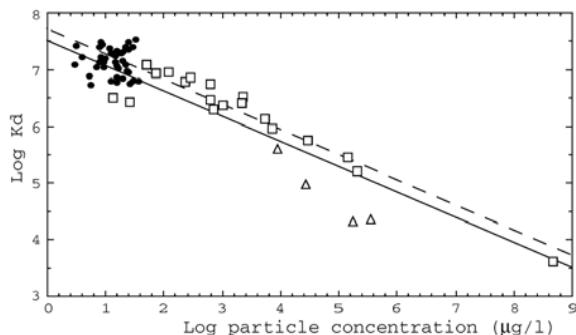
K_d = X/C_d = distribution coefficient



Henderson et al., 1999
based on Honeyman et al., 1988

Particle concentration effect

Effect of reduction in Kd on scavenging in BNL

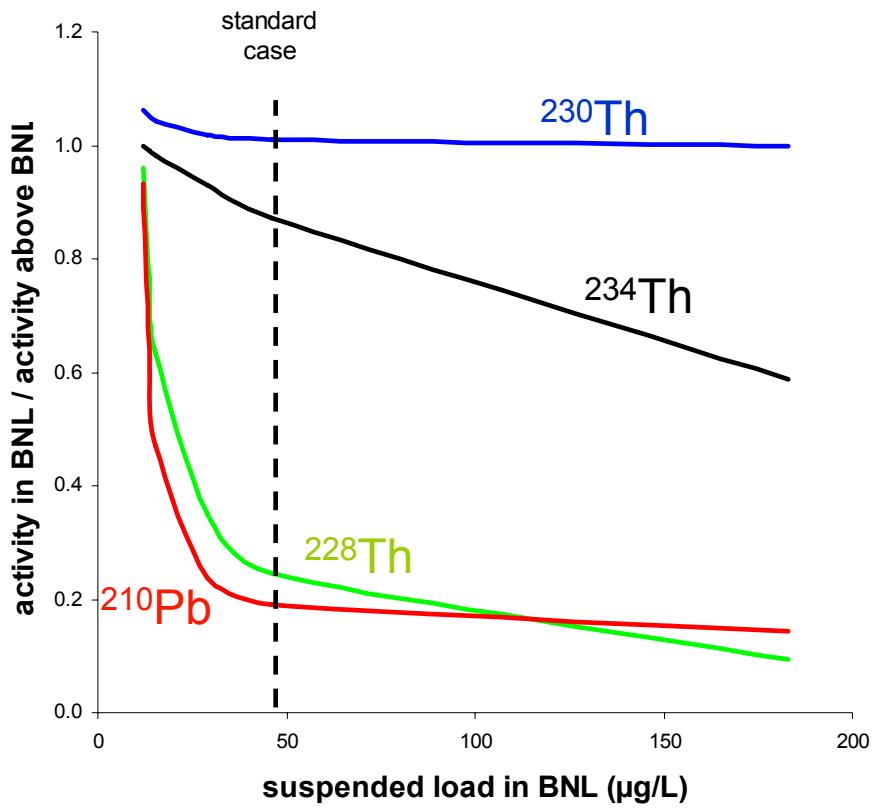


Half-lives: ^{234}Th : 0.07y, ^{228}Th : 1.9y; ^{210}Pb : 22.3y; ^{230}Th : 75400y

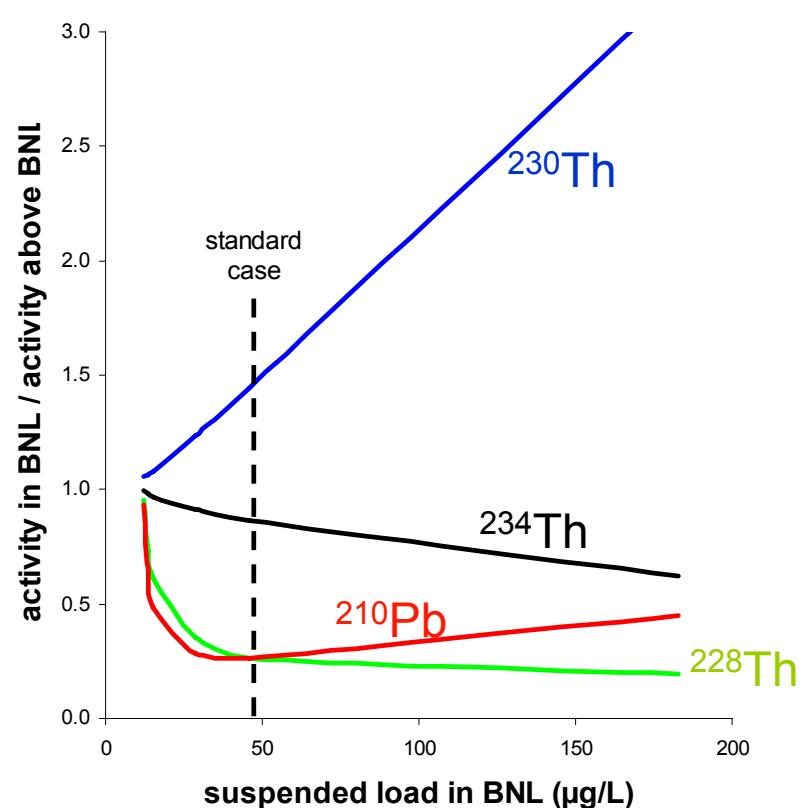
Resuspension

Effect on scavenging in BNL

dissolved activity



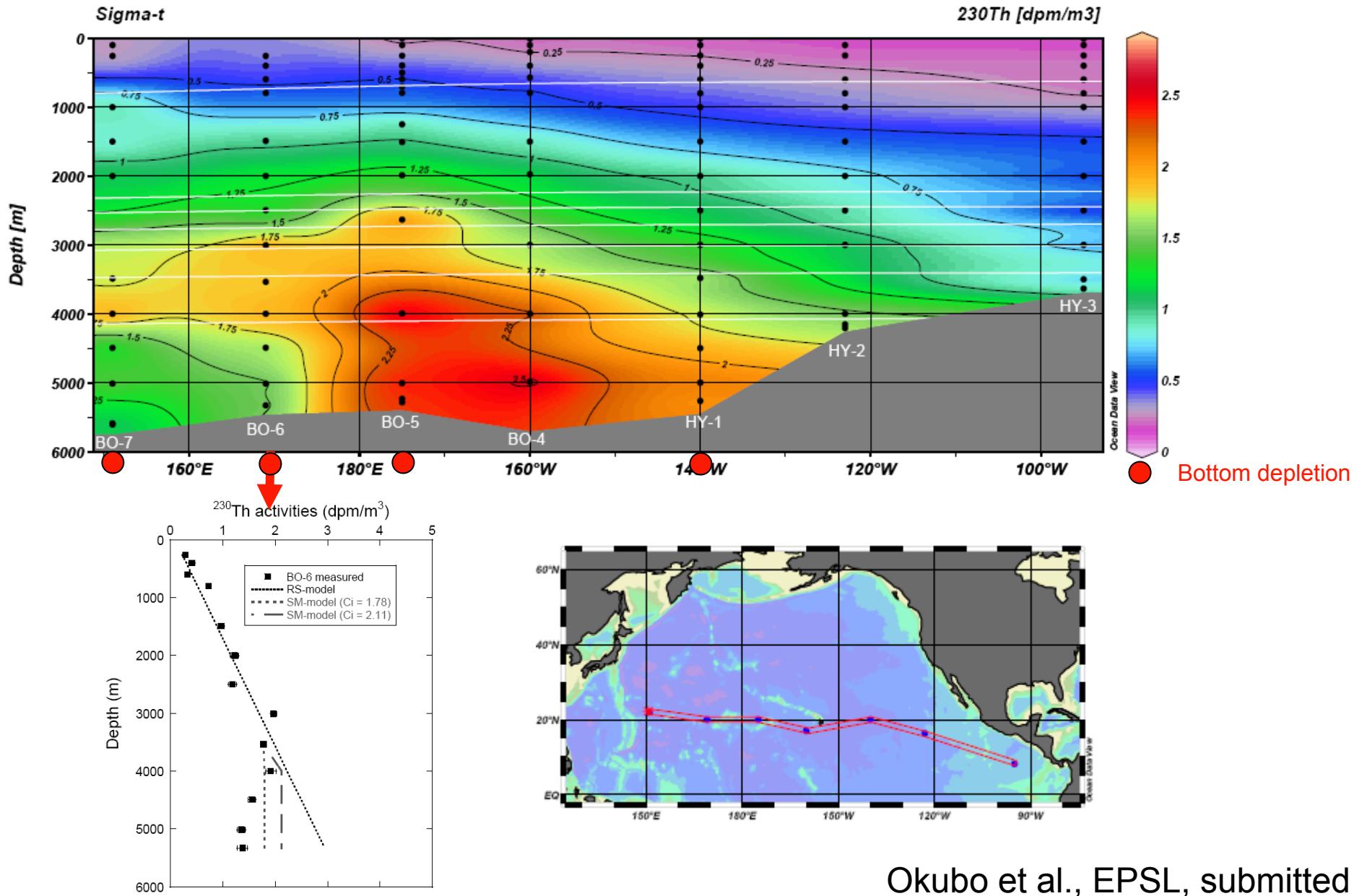
total activity



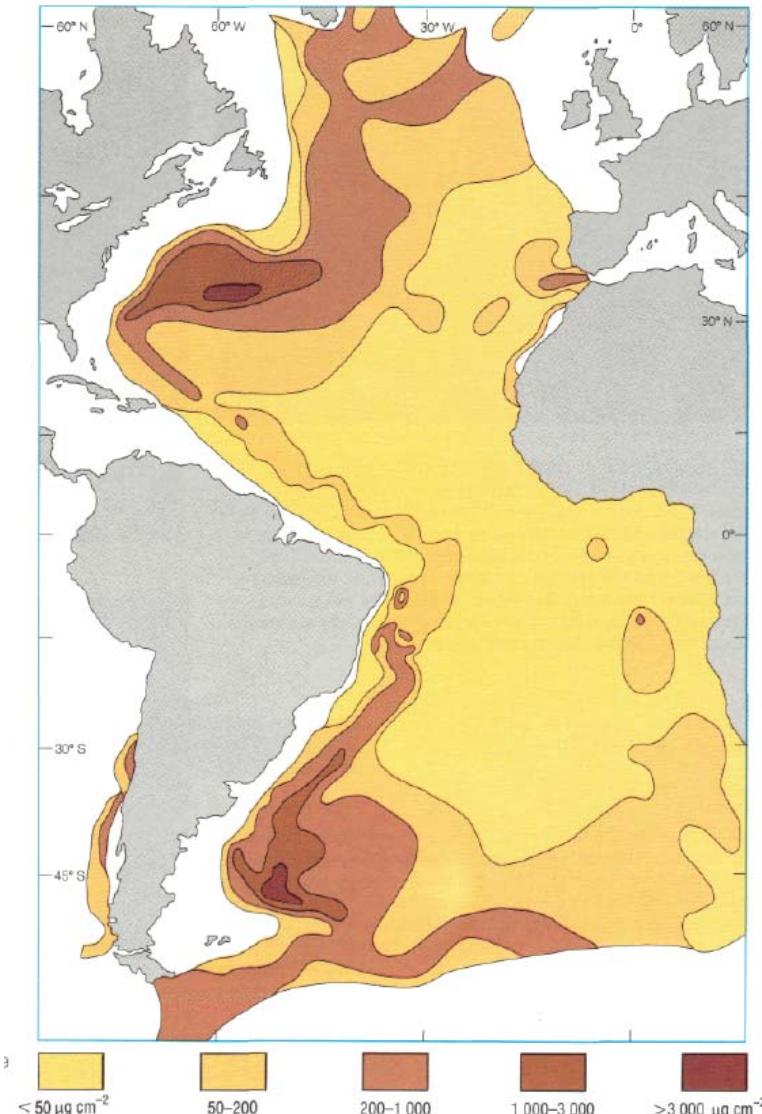
Conclusions on local resuspension

- Interaction between **resuspension** and **bioturbation** causes changes in dissolved component in BNL (scavenging)
 - if tracer decays within bioturbated zone
 - if Kd changes as a result of
 - Diagenetic changes (eg MnO_2 enrichment)
 - Particle dynamics (particle concentration effect, aggregation/disaggregation)

^{230}Th section in Pacific



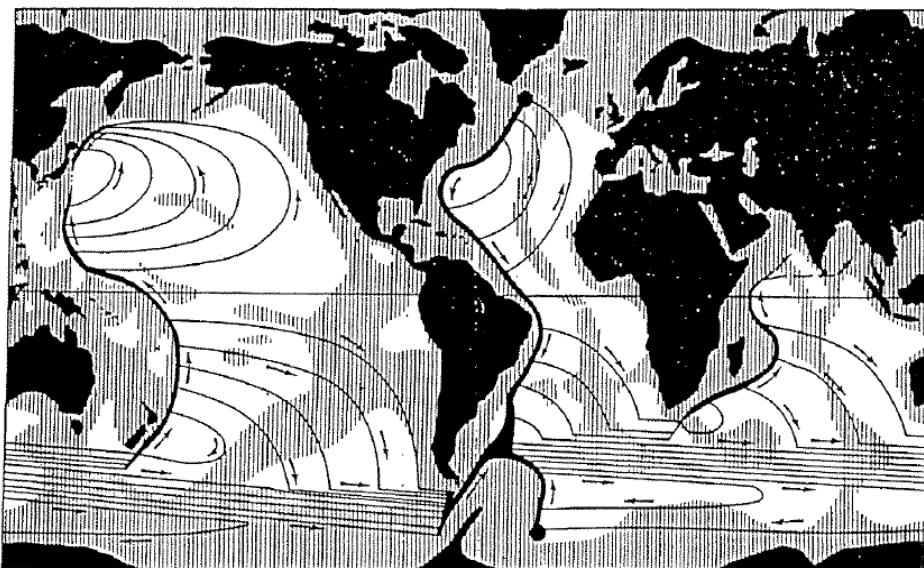
Long-range transport in BNL



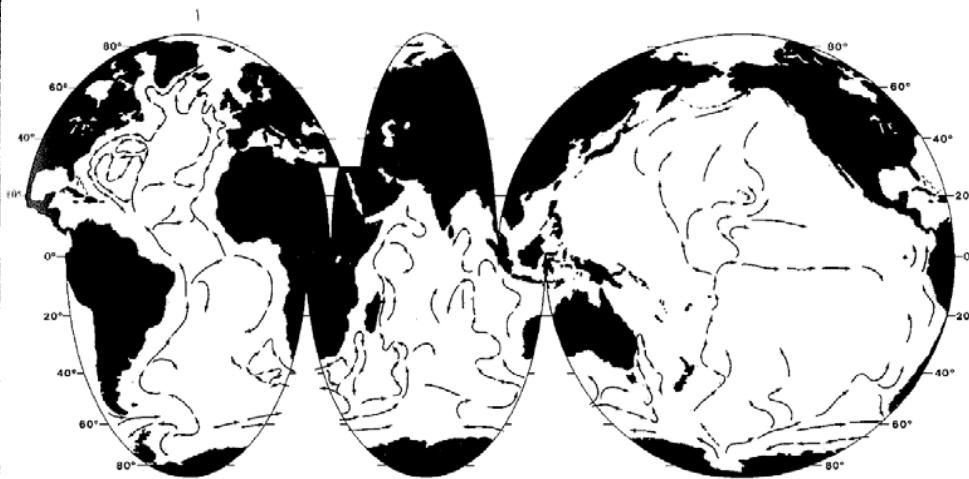
Suspended Particle Inventory in BNL
(in excess of clear water minimum)

Open University based on
Biscaye and Eittreim 1977

Deep Western Boundary Currents

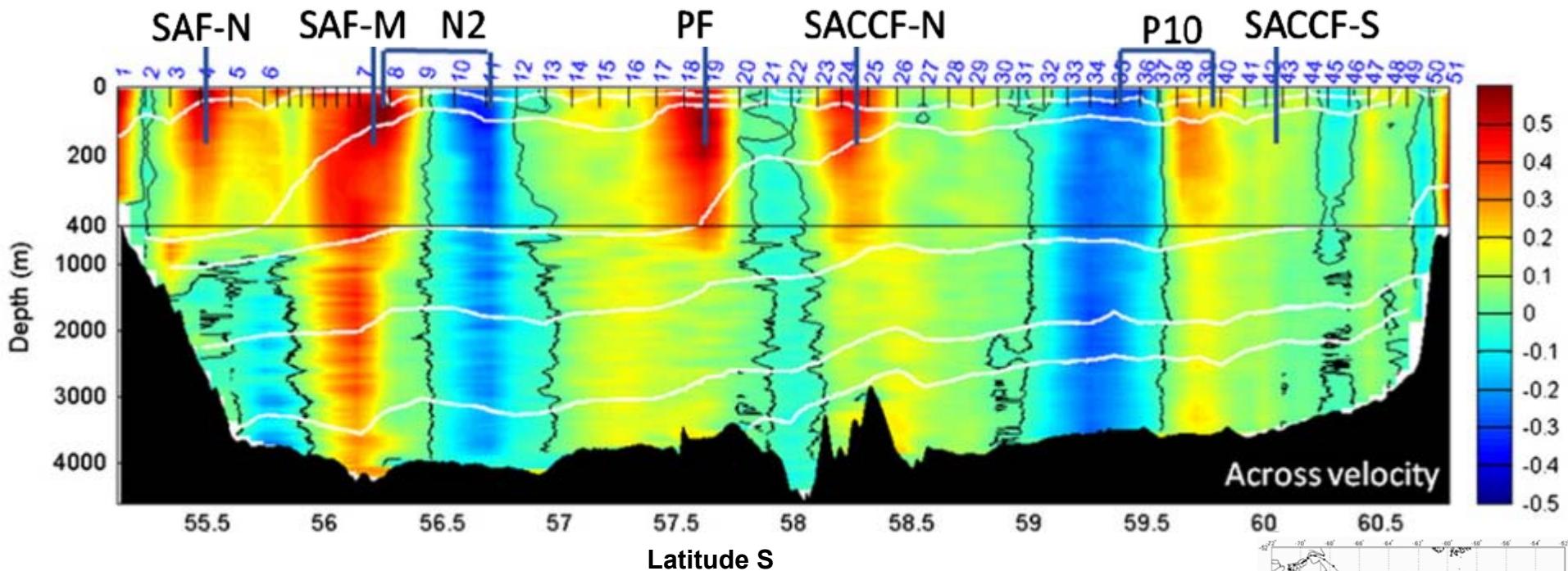


Stommel, 1958

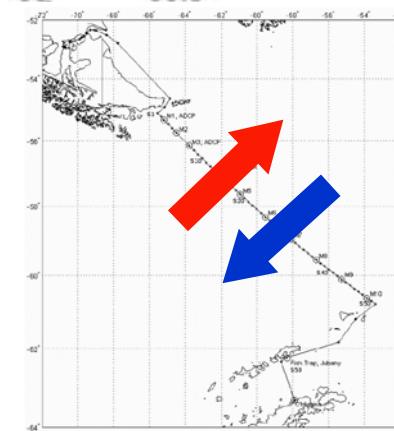


McCave 1986

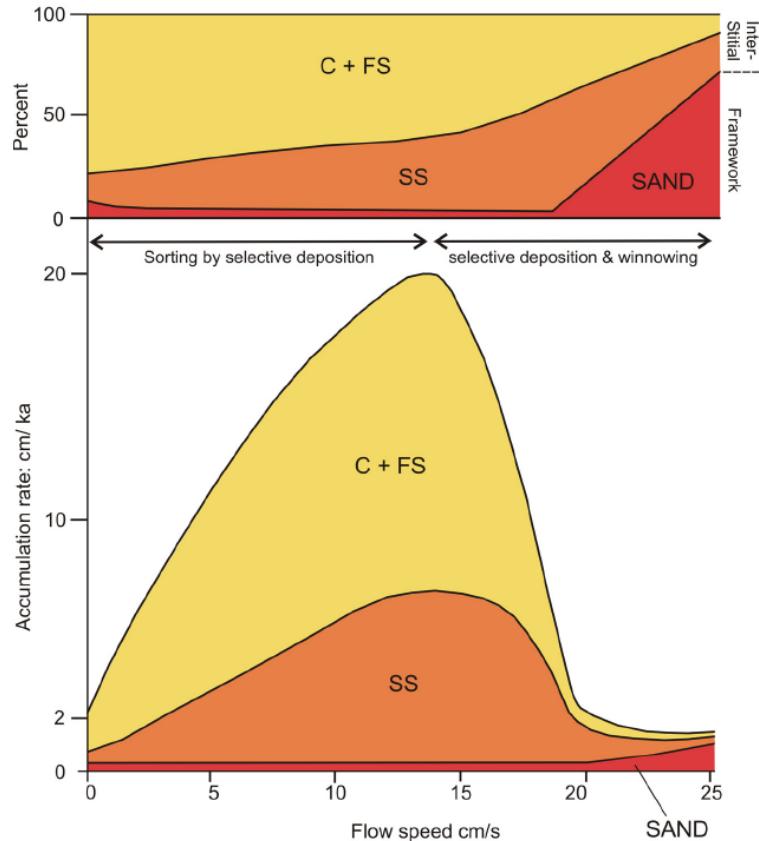
Currents through Drake Passage



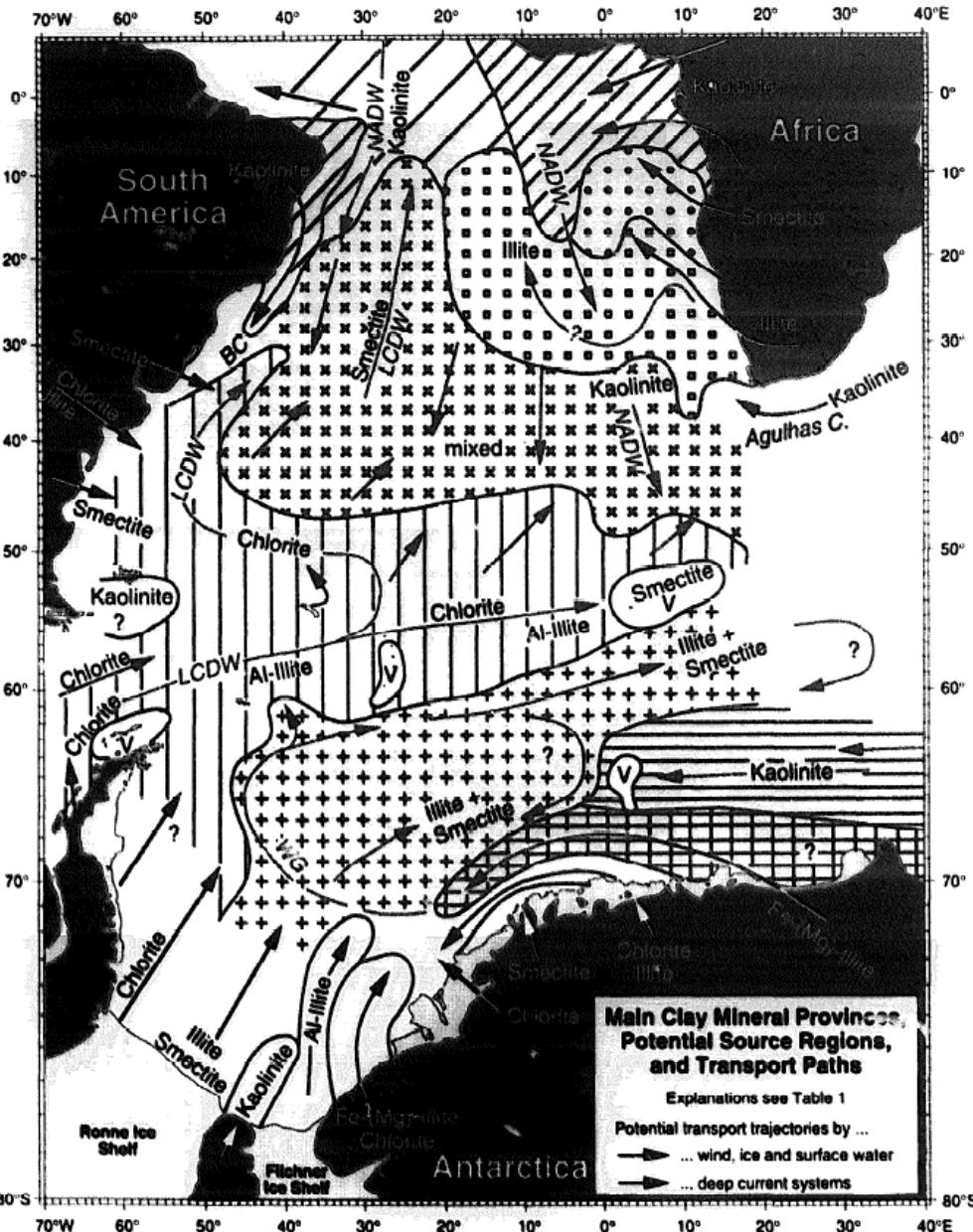
LADCP-velocity (m/s)



Resuspension and sorting in BNL

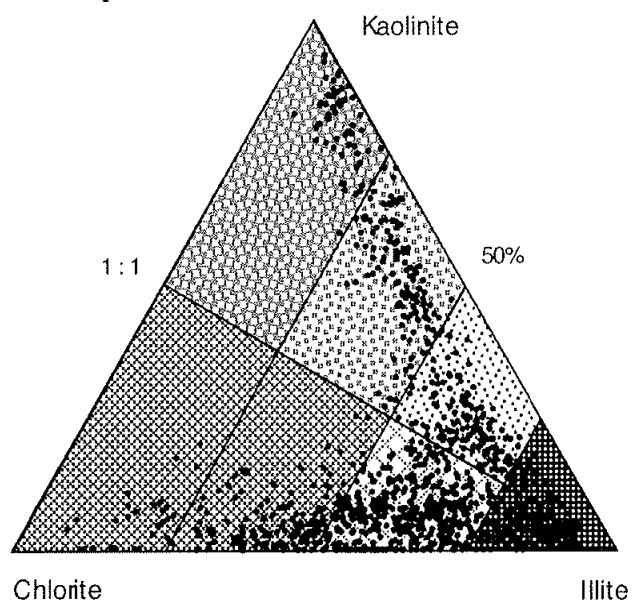
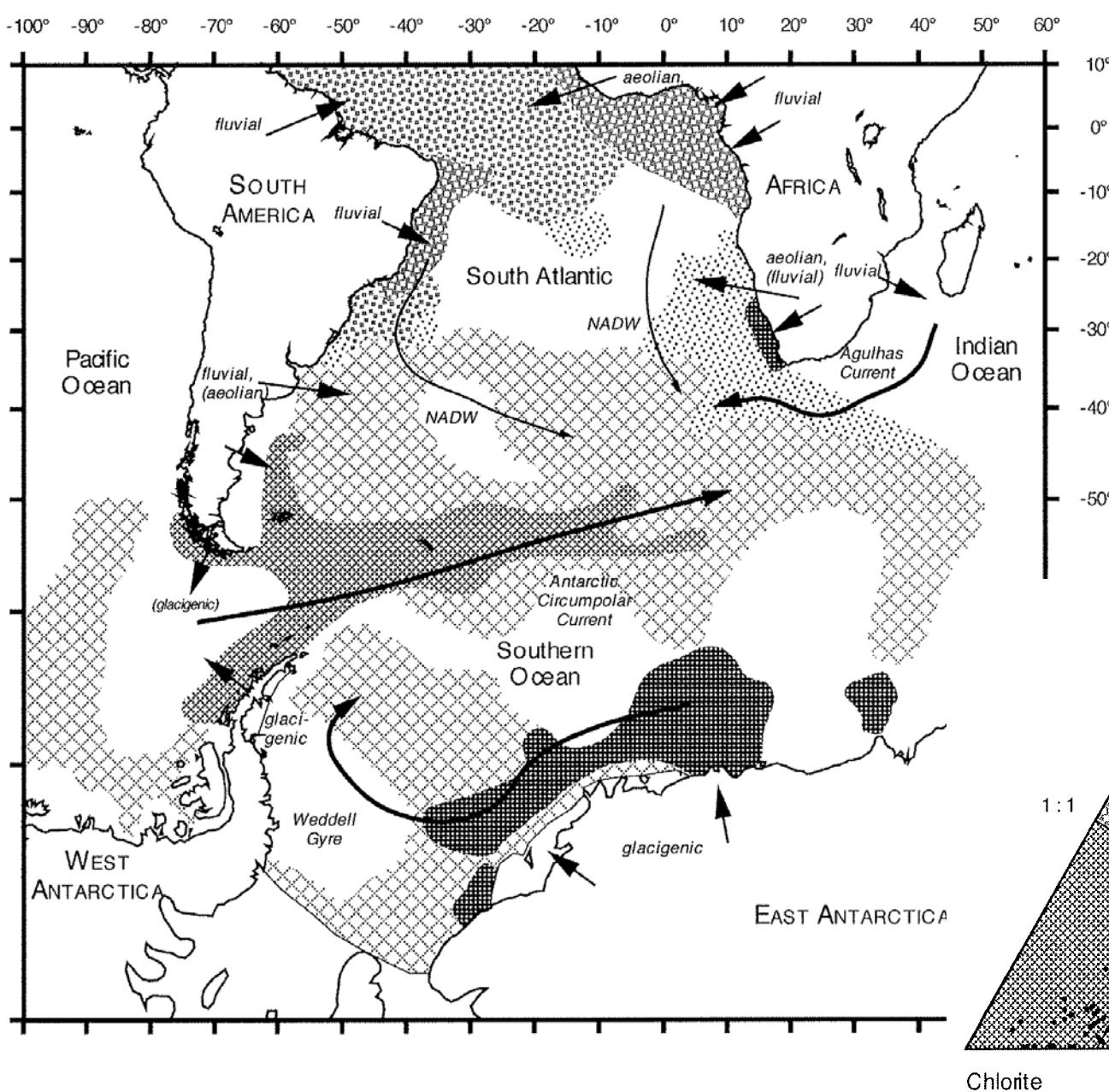


C+FS, clay and fine silt ($<10\text{ }\mu\text{m}$)
 SS, sortable silt (10–63 μm)
 SAND ($>63\text{ }\mu\text{m}$)



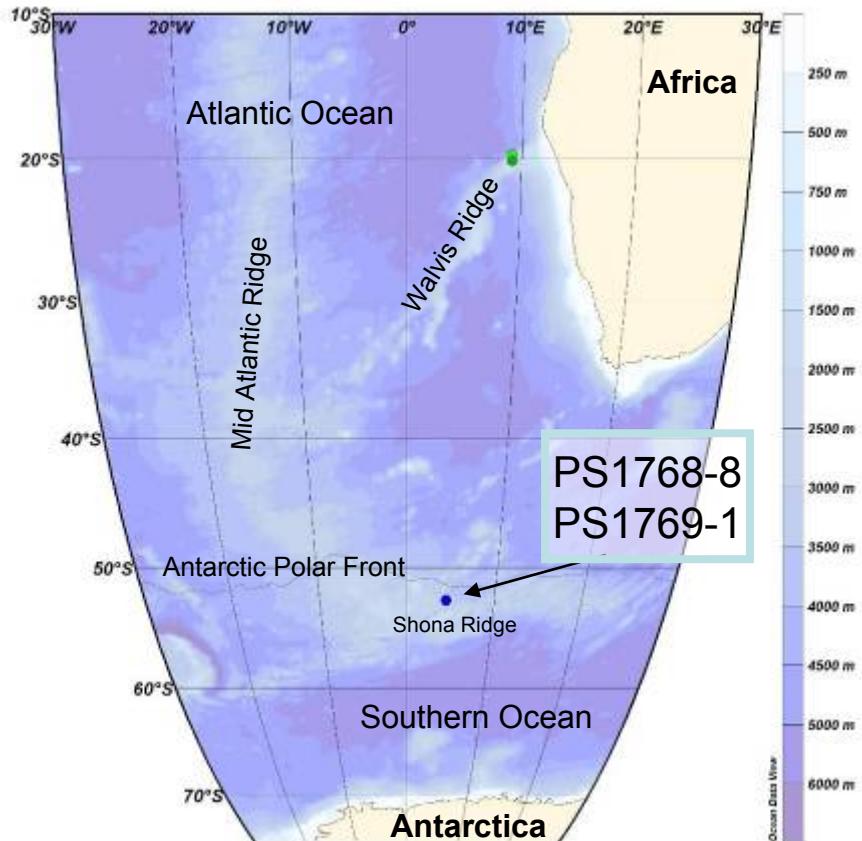
Petschik et al., 1996

Clay mineral transport



Diekmann et al., 2004

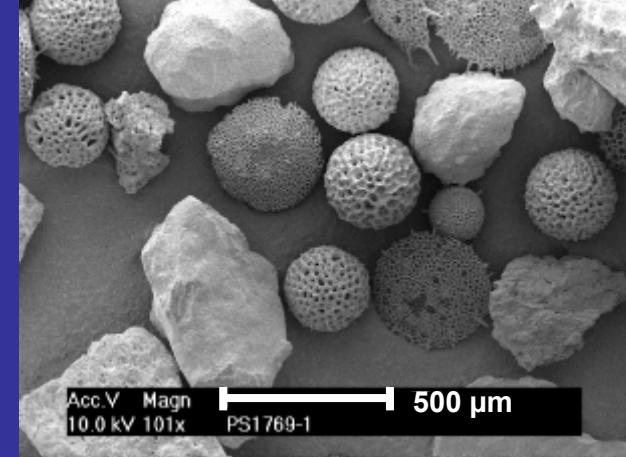
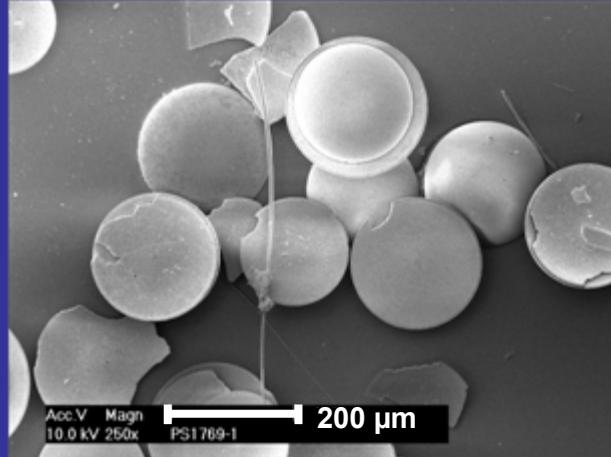
What is the effect of sorting on Transport of TEI on particles ?



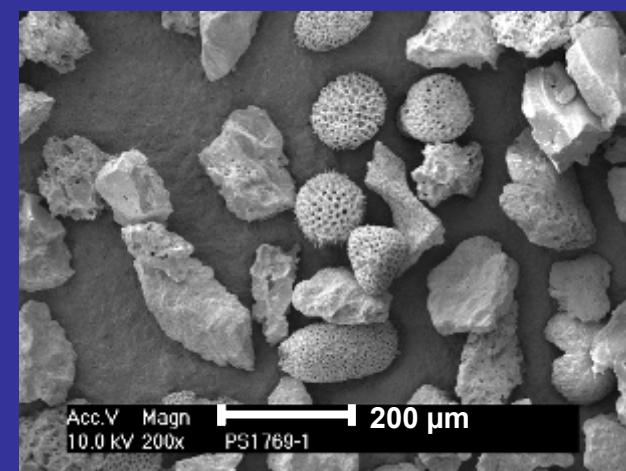
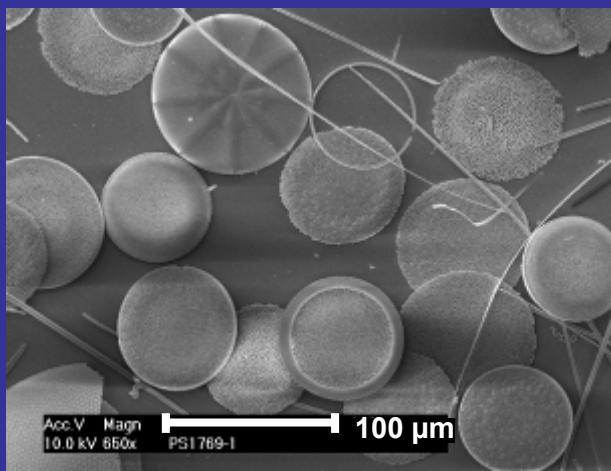
Cores studied by Kretschmer

PS1769-1
16 ka

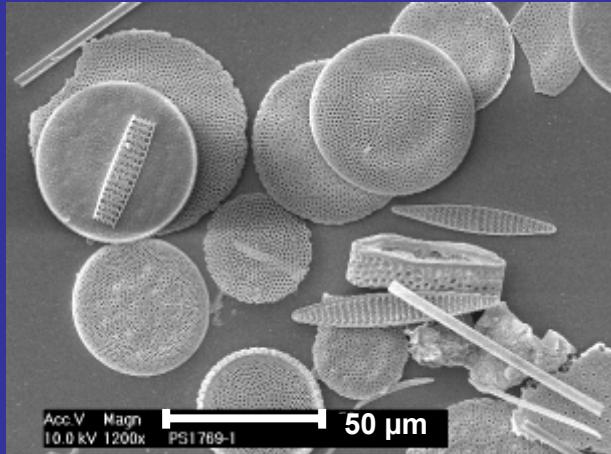
> 125 μ m
slowly settling



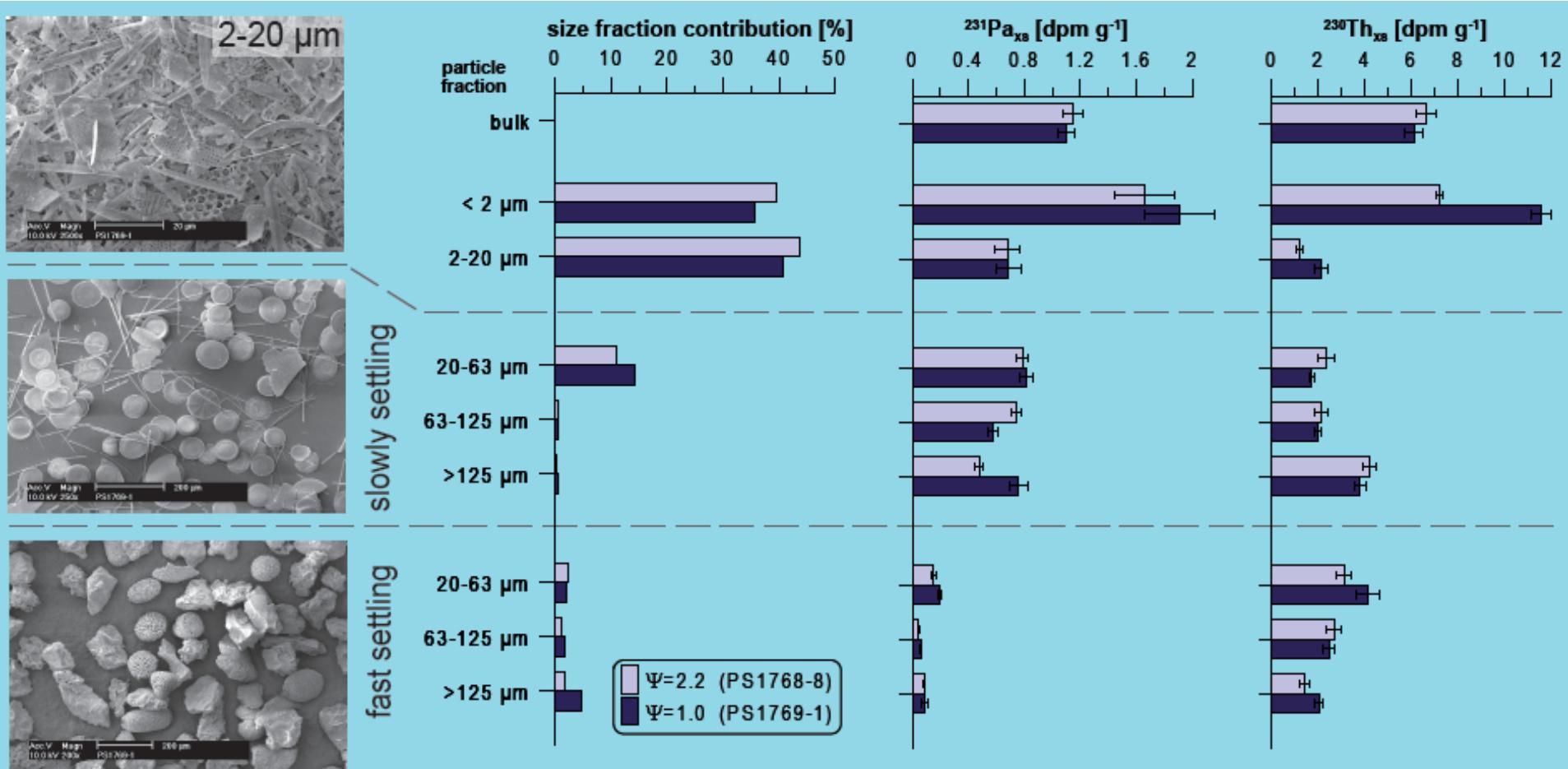
63-125 μ m
slowly settling



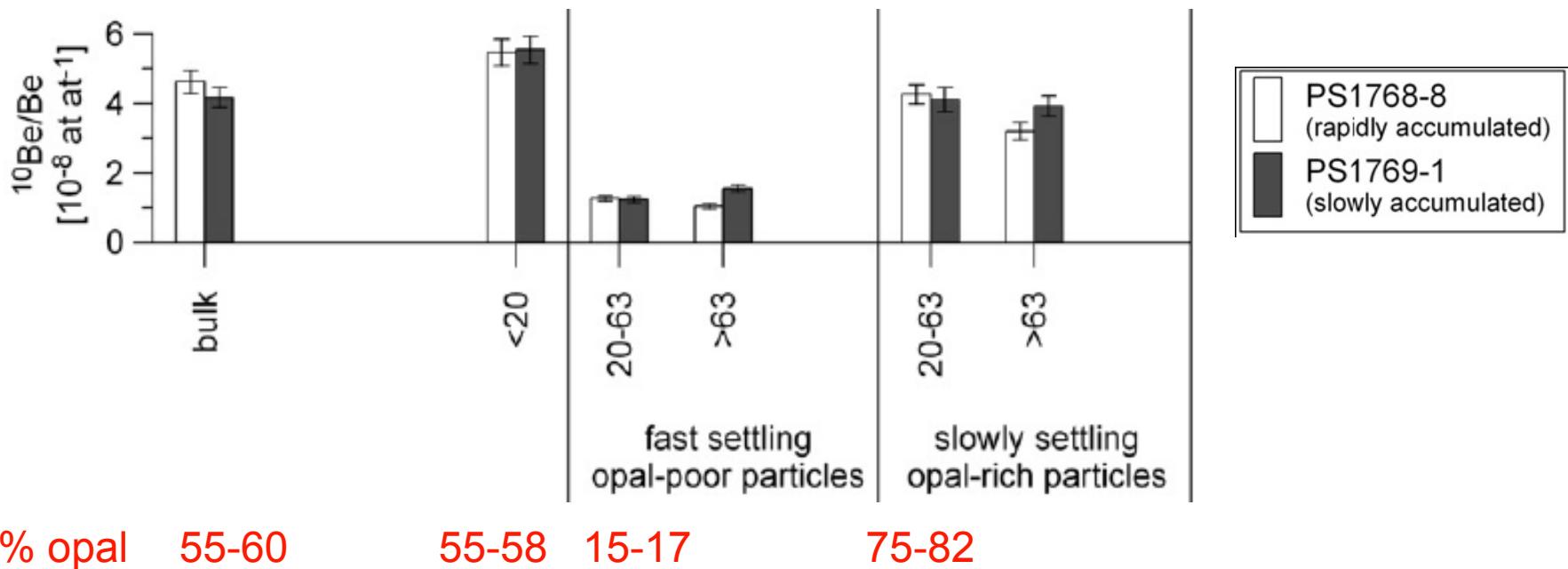
20-63 μ m
slowly settling



^{230}Th and ^{231}Pa fractionation by size and sinking rate

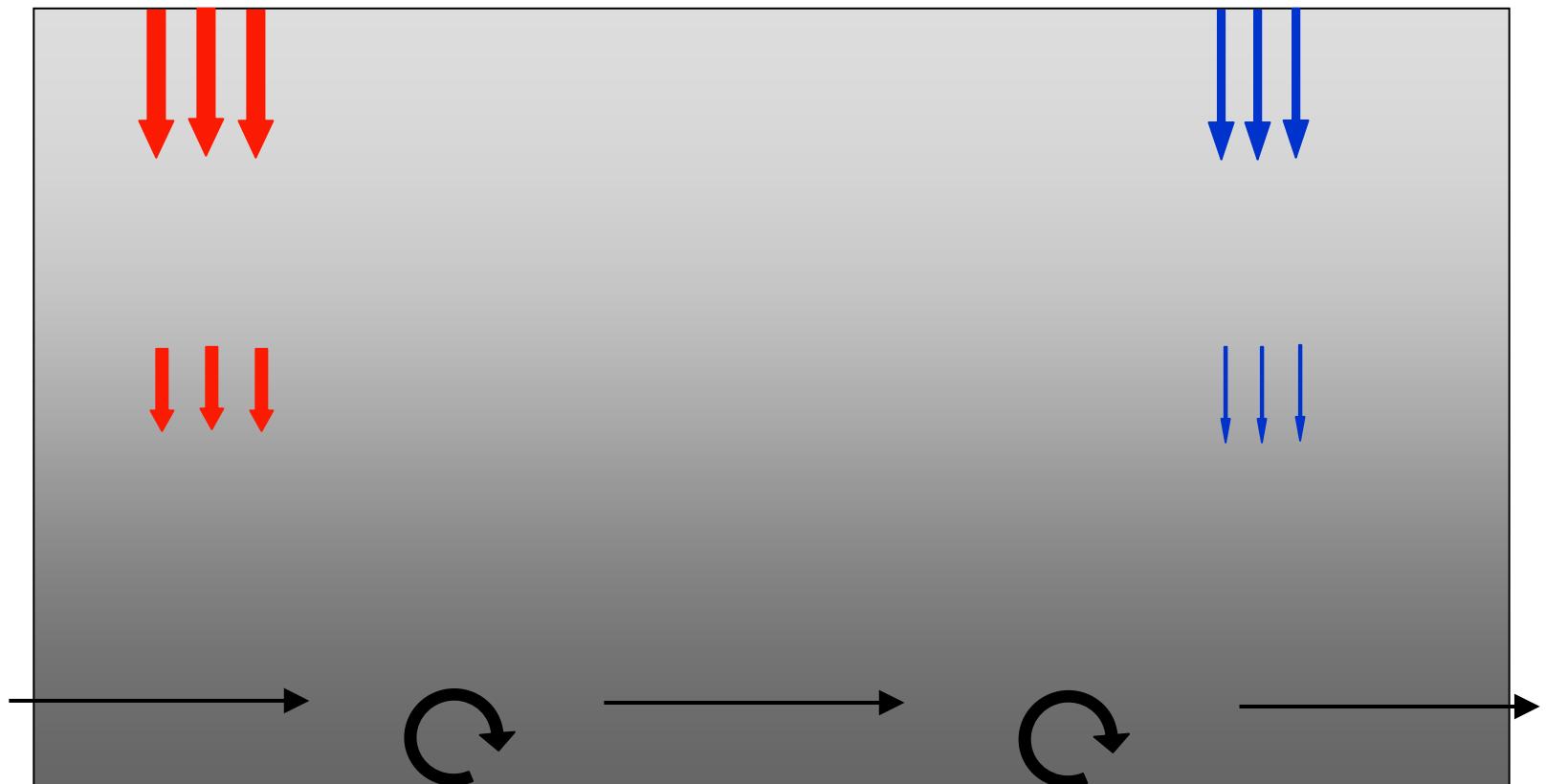


Evidence for remote input of ^{10}Be



Lithogenic part of $<20\mu$ fraction enriched in ^{10}Be from deep water

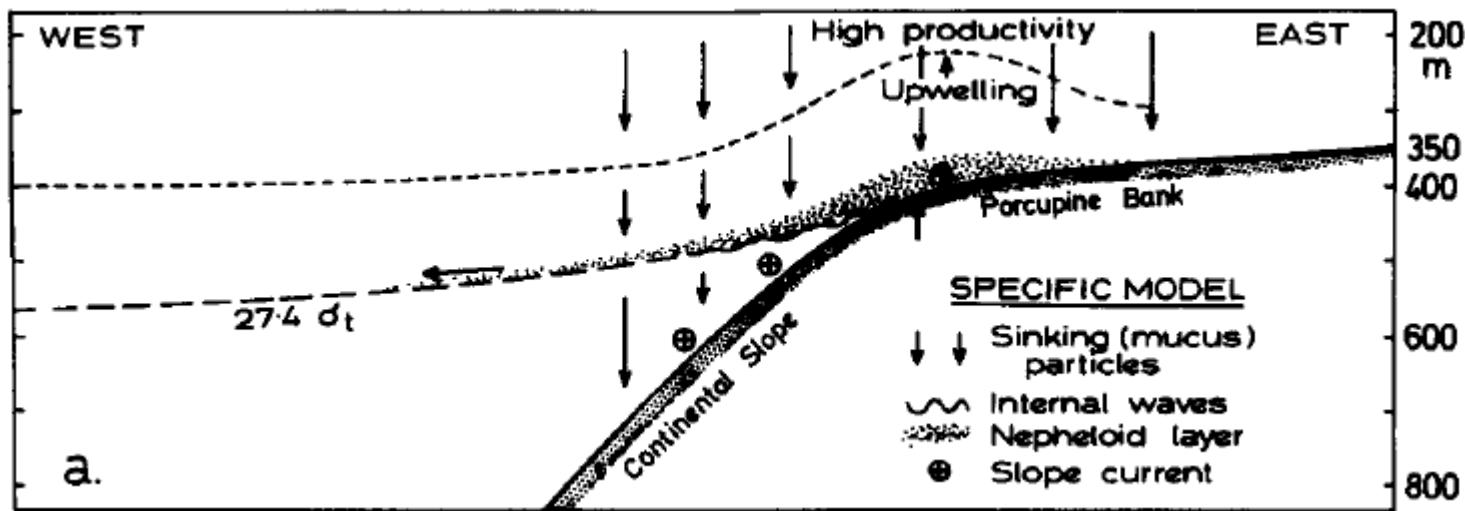
- 1- Particles sink through advected water masses and reequilibrate
(sinking particles acquire isotopic signature of deepest approx 1000m)
(^{231}Pa : Thomas et al., 2006)
- 2- a slowly settling fraction of particles (clay, opal) is transported over large distances in the BNL where they mix and exchange with locally produced particles (conflicts with assumptions of ^{230}Th normalization)

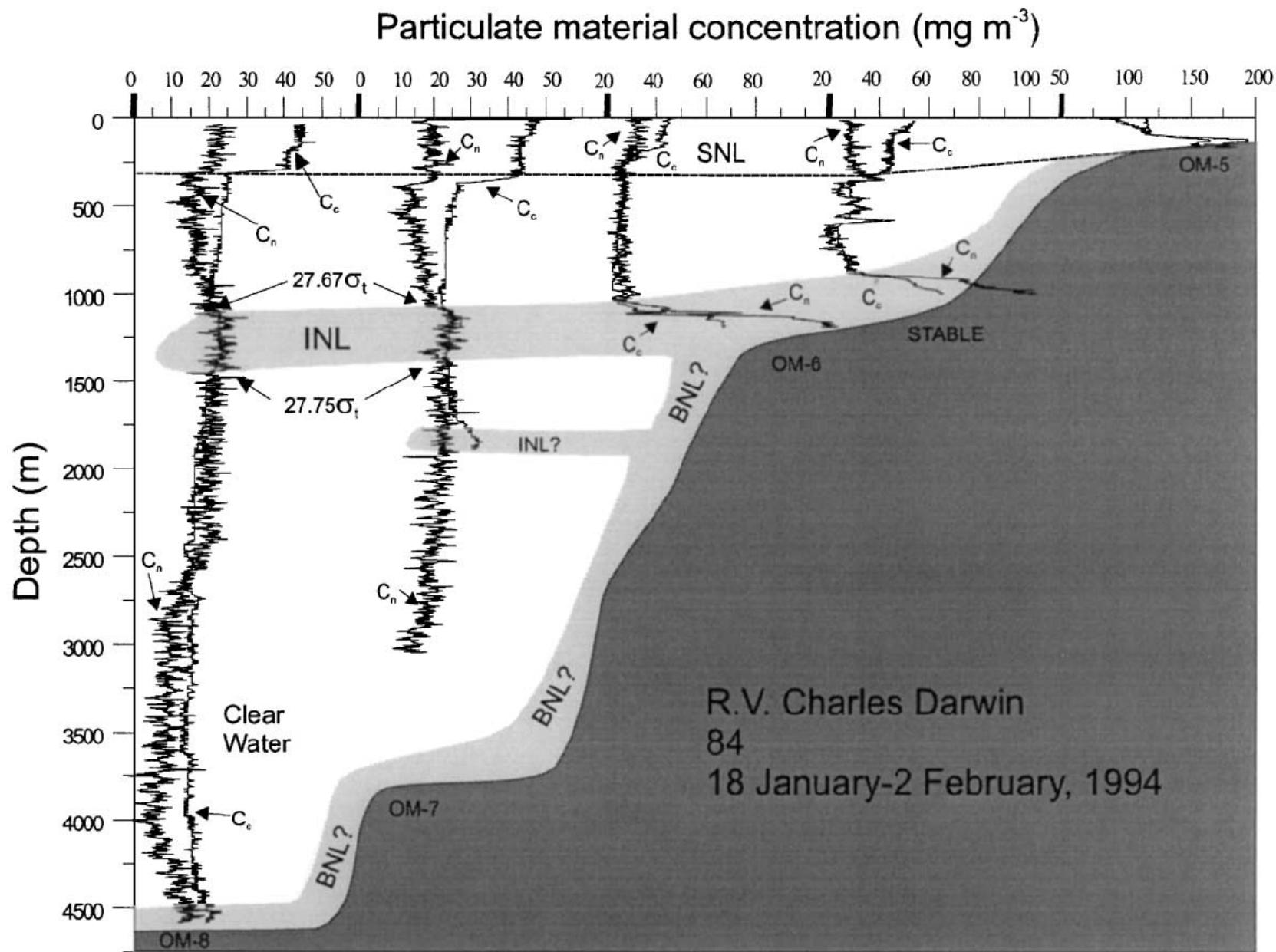


Conclusions on long-range transport in BNL

- Transport among areas with widely different primary sedimentation
- Composition of material suspended in BNL is different from that in clear water above it (cf data Rob Sherrell, Lars Stemmann; what about organic coatings: Peter Santschi?)
- Grain size fractionation. Long range transport of clay minerals
- ^{230}Th , ^{231}Pa and ^{10}Be adsorb preferentially onto the smallest grain sizes
- $^{231}\text{Pa}/^{230}\text{Th}$ and $^{10}\text{Be}/^{230}\text{Th}$ is enhanced in slowly settling pure opal fraction
- Settling rate fractionation during focussing causes increase in bulk ^{230}Th concentration and in $^{231}\text{Pa}/^{230}\text{Th}$ ratio

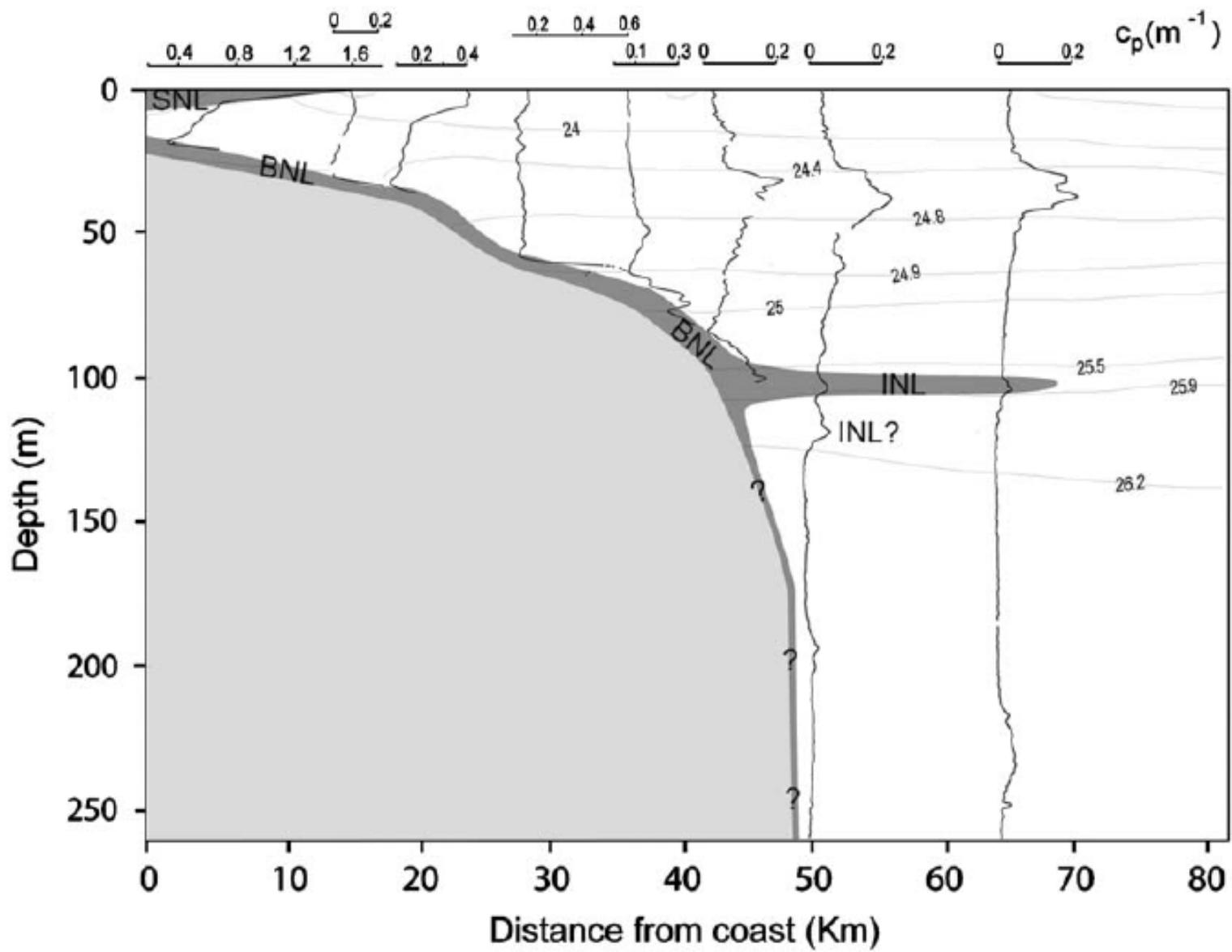
Intermediate Nepheloid Layers





OMEX: European margin at 47°-50°N

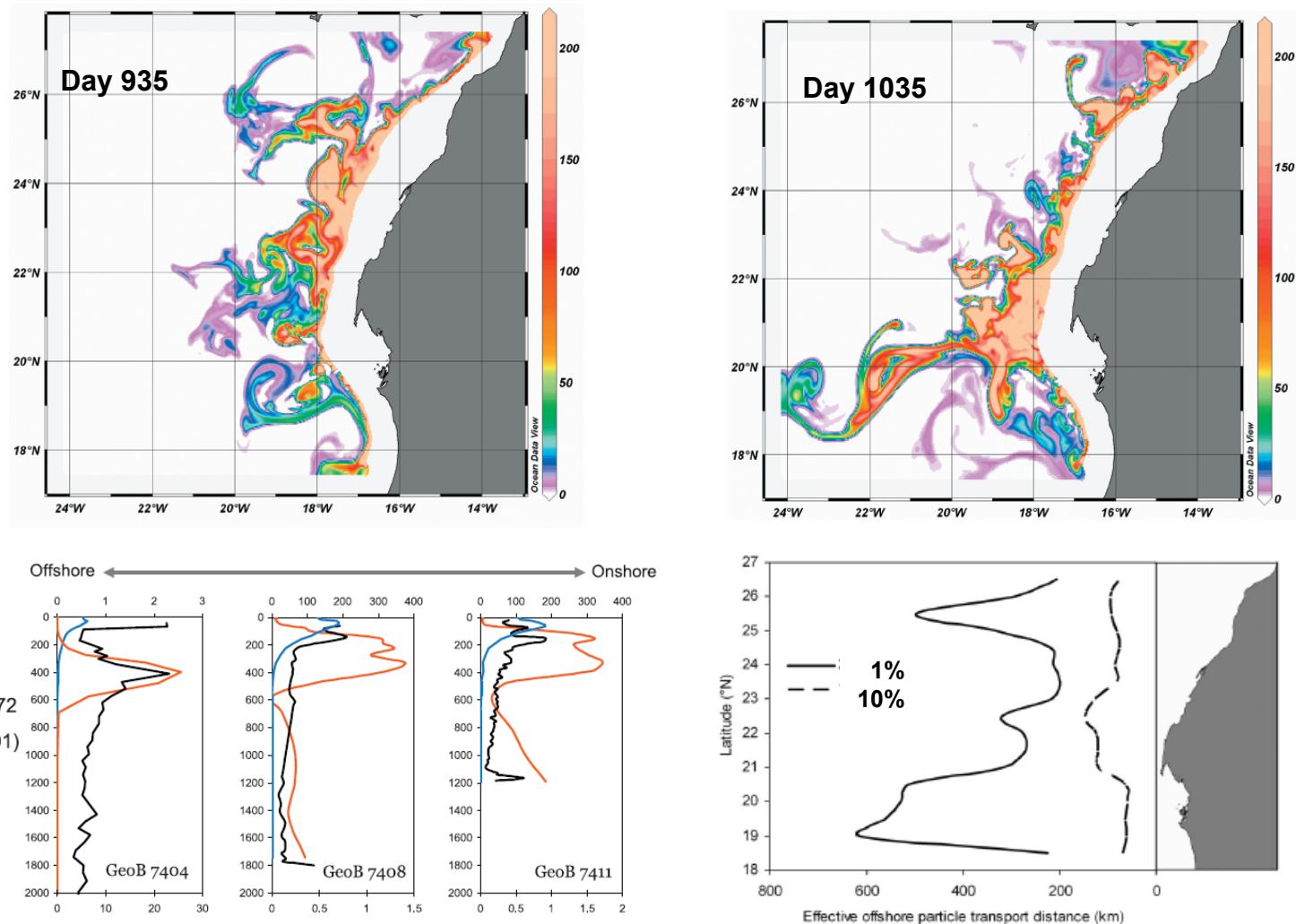
McCave et al., 2001



Cariaco Basin, Venezuela

Lorenzoni et al., 2009

Modelled distribution at 500m depth of shelf-derived particles with settling rate of 5 m d⁻¹

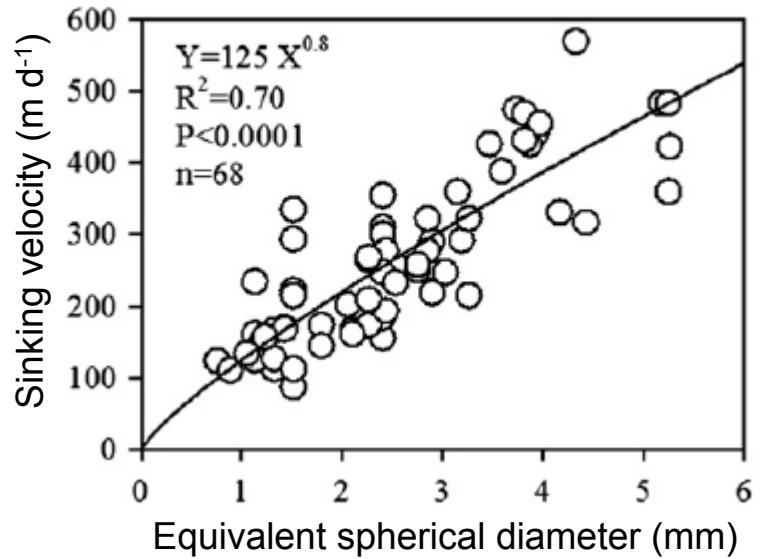
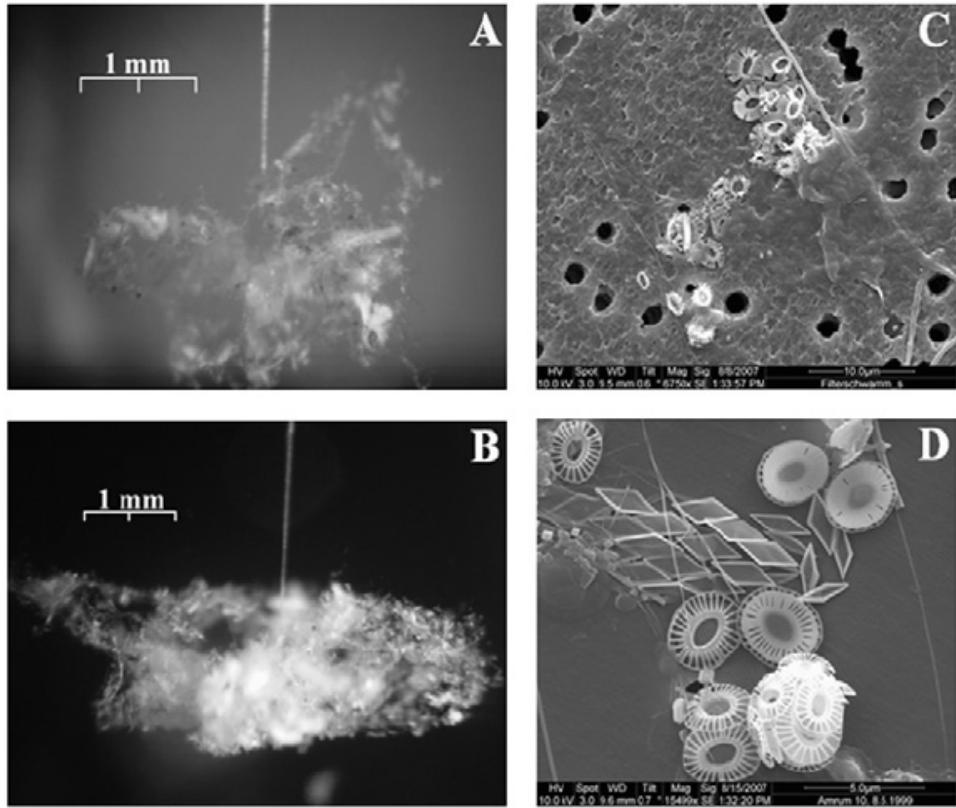


West Africa

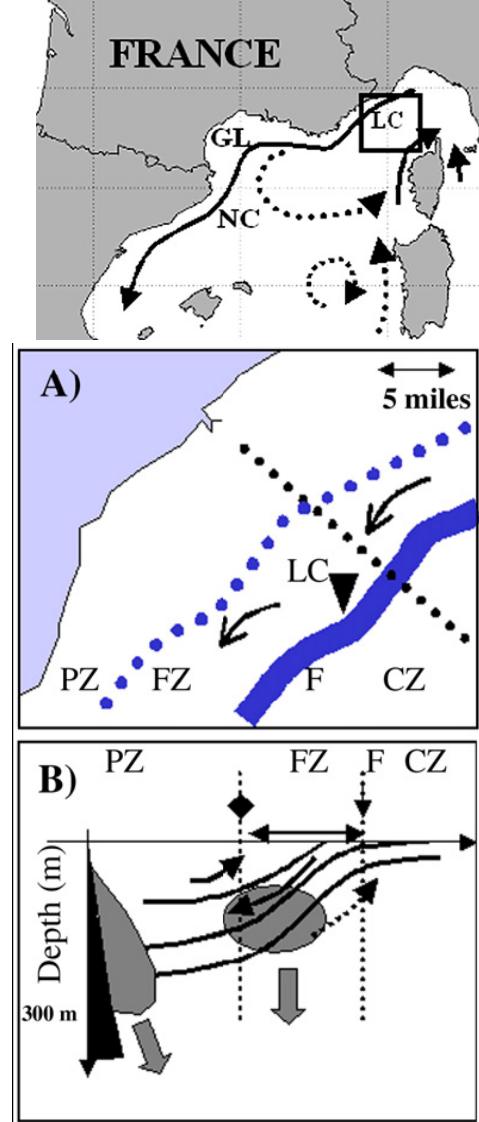
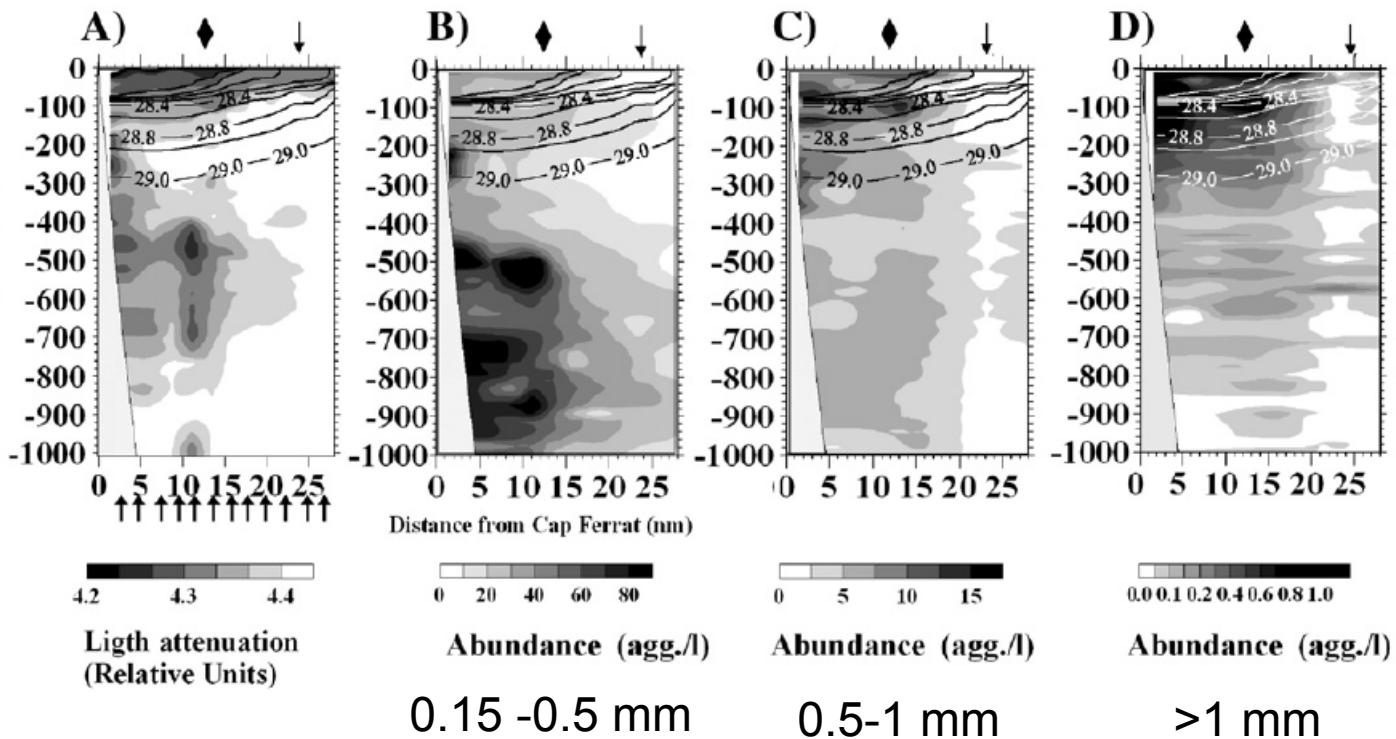
Karakas et al., 2006

Aggregation and sinking rates in INL

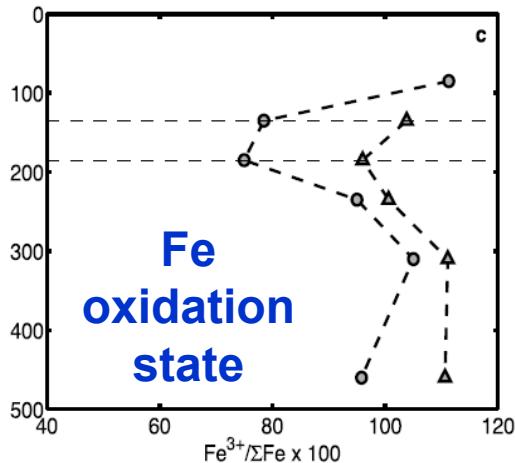
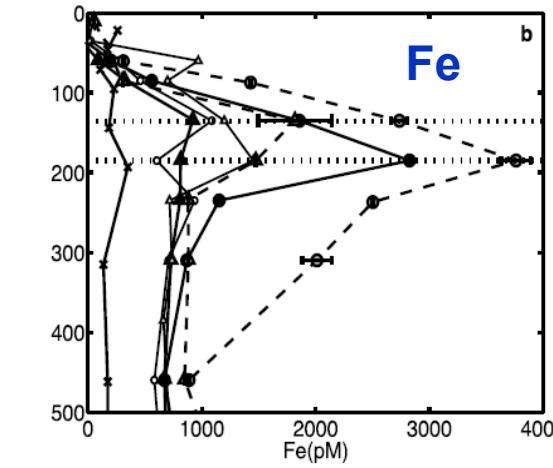
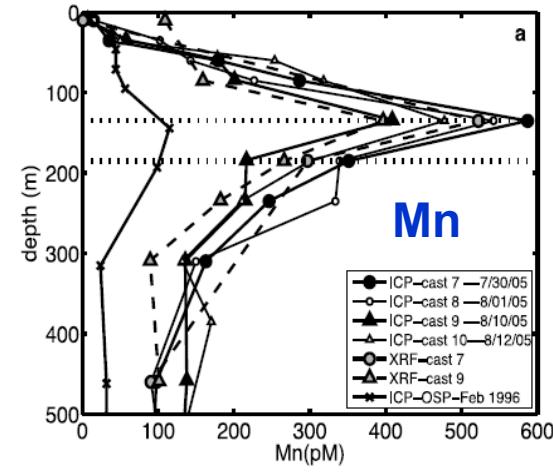
Aggregates formed in roller tanks
In samples from fluorescence max.



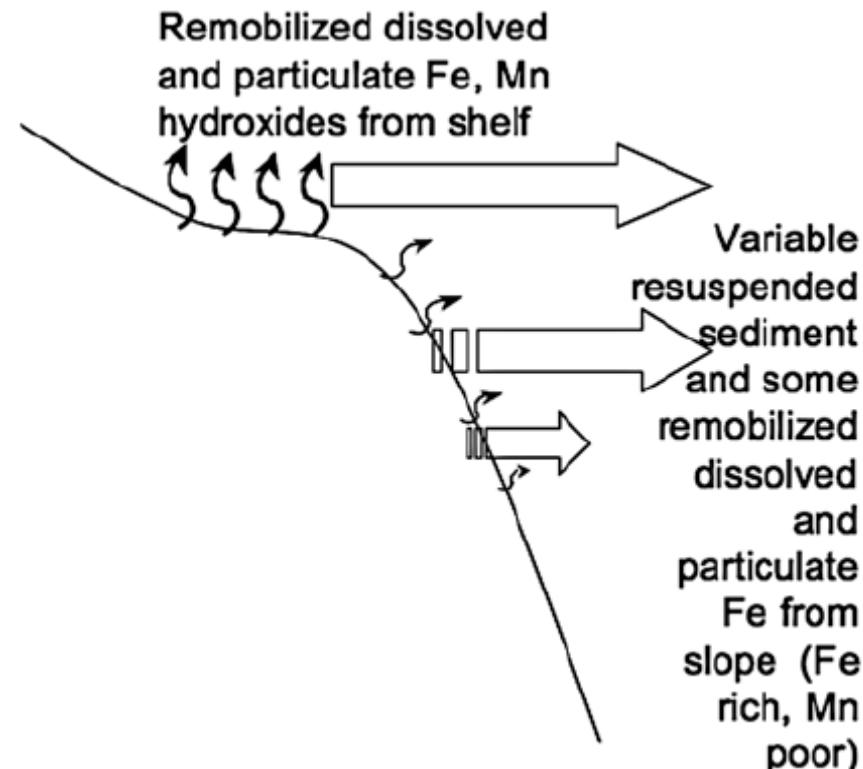
Depth (m)



Particulate (1-51 μ m) Mn and Fe In NW Pacific



NW Pacific



Issues on TEIs in INL

- Link between dispersal of particulate (INL) and dissolved tracer signals (e.g. Fe, Mn, ^{210}Pb , Nd)
- Time scale of distribution of shelf inputs (Ra isotopes, ^{228}Th)
- Is the lateral distribution of shelf inputs limited by aggregate formation?

Approach for better representation of processes in nepheloid layers

Measurements:

- Determine adsorption equilibrium in BNL (Pa, Th, Nd, Be)
 - This requires measurement of particle composition and mass concentration
- Determine aggregation and settling rates
- Follow particulate and dissolved signals in INL.

Models:

- Include several particle classes with different adsorption characteristics and settling rates
- Include resuspension/selective deposition to allow for long range transport in BNL

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Parameter values resuspension model

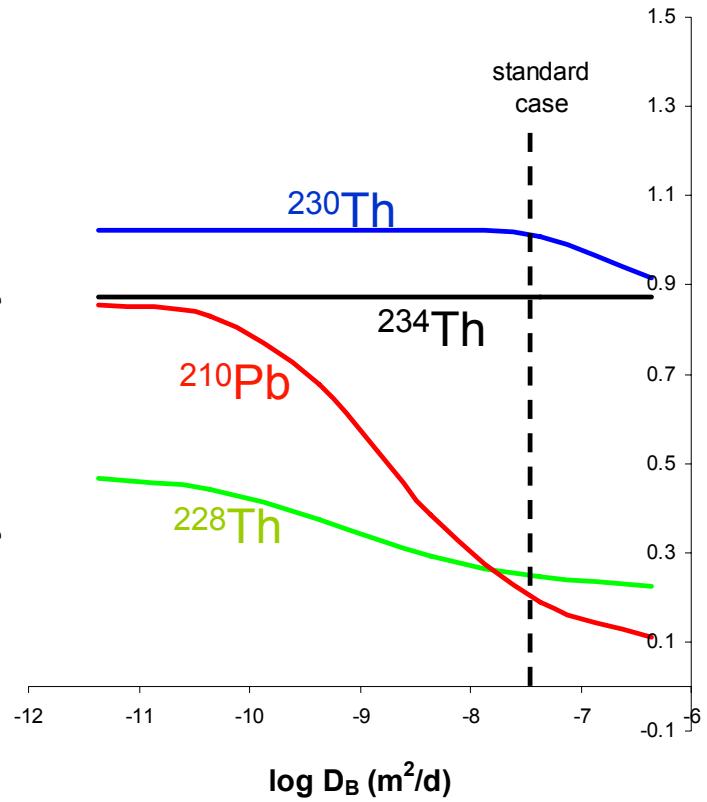
DB	4.3E-08	bioturbation coefficient	$m^2 d^{-1}$	$5.0E-08 \text{ cm}^2 \text{s}^{-1}$
DM	8.6E-06	molecular diffusion coefficient	$m^2 d^{-1}$	$1.0E-06 \text{ cm}^2 \text{s}^{-1}$
E	86.4	eddy diffusion coefficient in BNL	$m^2 d^{-1}$	
h	100	Height of bottom boundary layer (BNL)	m	
k1c	0.001479	adsorption rate constant in clear water	d^{-1}	
k-1	0.008219	desorption rate constant	d^{-1}	
K	0.01	resuspension rate	d^{-1}	
Kd	8.5E+06	distribution coefficient	- (vol/vol)	
L	5.0E-04	Thickness of resuspension zone (surface sediment)	m	
w	-10		cm kyr^{-1}	
ws	-5.8	settling rate	md^{-1}	
alpha	1	reciprocal of characteristic resuspension height	m^{-1}	
phi0	0.9	liquid volume fraction in sediment (porosity)	-	

Bioturbation

Effect on scavenging in BNL

dissolved activity

activity in BNL / activity above BNL



total activity

activity in BNL / activity above BNL

